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ABSTRACT

Various plate reconstructions predict that the Alexander terrane, a Neoproterozoic–Jurassic crustal fragment now located in the North American Cordillera, evolved in proximity to the northern Appalachian-Caledonian convergent margin during assembly of supercontinent Laurussia. To test stratigraphic connections with Laurussia that are implied by these plate reconstructions, we measured the Hf isotopic compositions of 176 detrital zircons from two relevant sedimentary sequences of the Alexander terrane. An older, Upper Silurian–Lower Devonian terrestrial to shallow-marine molasse sequence yields 405–490 Ma detrital zircons with negative $\varepsilon_{Hf(t)}$ values and Mesoproterozoic to Paleoproterozoic Hf model ages. In combination with paleomagnetic and biogeographic constraints, these Hf data argue for the molasse strata to be now-displaced equivalents of the Old Red Sandstone and primarily sourced from crustally contaminated granitoids in the Greenland, Svalbard, or British Caledonides. Late Silurian–Early Devonian orogenesis in the Alexander terrane is therefore likely related to the Scandian-Salinic phase of Appalachian-Caledonian mountain building. Younger, Middle Devonian sequences of the Alexander terrane are endowed in 390–490 Ma detrital zircons with positive $\varepsilon_{Hf(t)}$ values and Neoproterozoic Hf model ages. These isotopic signatures are consistent with the erosion of local basement rocks during the opening of the Slide Mountain–Angayucham backarc rift and tectonic separation of the Alexander terrane from northern Laurussia.

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INTRODUCTION

Earth's geography and plate-tectonic history during the early Paleozoic were dominated by the complex closure of the Iapetus Ocean, which involved multiple subduction zones and culminated with the Scandian-Salinic collision between greater Baltica (Baltica + Ganderia/Avalonia) and Laurentia (van Staal et al., 2009). Subduction of greater Baltica beneath Laurentia represents a major mountain-building event in the Appalachian-Caledonian orogen and led to the formation of supercontinent Laurussia (McKerrow et al., 2000). The main Scandian-Salinic phase of mountain building (400–440 Ma) was in part marked by the generation of crustally contaminated granitoids in Svalbard, Greenland, the British Isles, and Atlantic Canada (e.g., Whalen et al., 2006; Augland et al., 2012). Upper Silurian to Upper Devonian molasse strata that underlie much of the North Atlantic region partially consist of red-bed clastic successions referred to as the Old Red Sandstone (Friend and Williams, 2000).

An unresolved problem relevant to the assembly of Laurussia is where the northern continuation of the Appalachian-Caledonian orogenic system extends into the present-day Arctic Ocean region (Pease, 2011). Some global plate reconstructions argue for the northernmost extent of the Appalachian-Caledonian convergent margin to have involved crustal fragments of Laurentian, Baltican, and/or Siberian affinity that are now located in the North American Cordillera and Canadian High Arctic (Colpron and Nelson, 2009, 2011; Cocks and Torsvik, 2011; Miller et al., 2011). The Alexander terrane (Figs. 1A and 1B) is a Neoproterozoic–Jurassic crustal fragment in the North American Cordillera that provides an opportunity to investigate these proposed northern Caledonian connections. A growing body of multidisciplinary evidence suggests that the Alexander terrane formed part of a Late Cambrian–Early Silurian arc system in proximity to the Scandinavian and Russian High Arctic regions of Baltica prior to being the site of Late Silurian–Early Devonian orogenesis and Old Red Sandstone–like red-bed sedimentation (e.g., Soja, 1994; Bazard et al., 1995; Gehrels et al., 1996; Butler et al., 1997; Soja and Krutikov, 2008; Beranek et al., 2012, 2013).

Despite the fact that published geochronological studies have shown red-bed molasse strata of the Alexander terrane to contain 420–450 Ma detrital zircons (Gehrels et al., 1996; Grove et al., 2008), these U-Pb age data by themselves are not capable of discriminating provenance from crustally contaminated granitoids of the northern Caledonides or juvenile arc rocks of the Alexander terrane. In an attempt to solve this problem and test the connections with northern Laurentia that are implied by some global plate reconstructions (e.g., Colpron and Nelson, 2009, 2011), we measured the Hf isotopic compositions of dated zircons from Silurian–Devonian sedimentary sequences of the Alexander terrane. Hf isotopes are robust geochemical tracers that identify the chemical maturity of the lithosphere from which zircons are derived and allow sediment provenance studies to distinguish juvenile (radiogenic) or evolved (unradiogenic) source components of similar age (Kinny and Maas, 2003).

BACKGROUND GEOLOGY

The Alexander terrane underlies 100,000 km² of the North American Cordillera in southeastern Alaska, the Saint Elias Mountains of eastern Alaska, southwestern Yukon, and northwestern British Columbia, and the



Figure 1. (A) Location map of the Alexander terrane. White circles indicate detrital zircon sample locations. (B) Circum-Arctic cratons, orogens, and locations modified from Colpron and Nelson (2011). AX-Alexander terrane; SV-Svalbard. (C) Stratigraphy of the Alexander terrane as compiled by Beranek et al. (2012). White circles indicate approximate stratigraphic locations of detrital zircon samples. Geological time scale is from Gradstein et al. (2004).

Coast Mountains of west-central British Columbia (Fig. 1A; Gehrels and Saleeby, 1987). The Alexander terrane is presently bounded by faults that record its mid-Mesozoic accretion to, and subsequent displacement along, the western margin of North America (e.g., McClelland et al., 1992).

Southeastern Alaska

The Karheen Formation in the Prince of Wales Island region of southeastern Alaska consists of Upper Silurian to Lower Devonian red-bed clastic, carbonate, and volcanic rocks that were deposited during the Klakas orogeny (Fig. 1C). This Late Silurian–Early Devonian tectonic event is broadly characterized by contractional deformation and greenschistfacies to locally amphibolite-facies metamorphism (Gehrels and Saleeby, 1987; Gehrels et al., 1987). The timing of the Klakas orogeny separates an earlier, Late Cambrian–Early Silurian convergent margin setting for the Alexander terrane from a later, Mississippian–Pennsylvanian phase of tectonic stability (Gehrels and Saleeby, 1987). Early Paleozoic (430–500 Ma) arc-type and syn- to post-Klakas (390–415 Ma) magmatic rocks in the Alexander terrane yield juvenile Nd-Sr (whole-rock) and Hf (zircon) isotopic compositions, suggesting little to no input from evolved crustal sources (Samson et al., 1989; Saleeby, 2000; Cecil et al., 2011; Beranek et al., 2012, 2013).

Paleomagnetic data require an Early Devonian magnetization for red beds of the Karheen Formation at $14^{\circ} \pm 4^{\circ}$ paleolatitude in the Northern or Southern Hemisphere (Bazard et al., 1995; Butler et al., 1997). Additional paleogeographic constraints on red-bed deposition are derived from fossil assemblages in surrounding units. Upper Silurian stromatolite reefs and mud mounds of the Heceta Formation (Fig. 1C) that crop out beneath the Karheen red beds were colonized by distinctive sphinctozoan sponges and hydroids (Soja, 1994). These biotas resemble those in the Uralian (east European) margin of Baltica, Salair region of southwest Siberia, and Siberian-affinity Farewell terrane of Alaska (Antoshkina and Soja, 2006). Early to Middle Devonian brachiopod, gastropod, and coral fauna in the Prince of Wales Island region also display statistical similarities with fossils of the Uralian margin, Siberian platform of southern Taimyr, and Kolyma-Omolon superterrane (Blodgett et al., 2002; Pedder, 2006). Larval exchange with the Baltican and Siberian realms during the Late Silurian-Middle Devonian strongly suggests a Northern Hemisphere position for the Alexander terrane within the Uralian Seaway, the marine corridor that separated Laurussia from Siberia (Soja and Antoshkina, 1997).

Saint Elias Mountains, Northwestern Canada

Terrestrial and shallow-marine strata of the lower Icefield assemblage in the Saint Elias Mountains are age equivalent to the Karheen Formation (Fig. 1C). Upper Silurian and Devonian rocks of the Icefield assemblage consist of red-brown to orange to gray weathering sandstone, conglomerate, shale, sandy limestone, and minor basalt to andesite flows and tuff (Mihalynuk et al., 1993). Evidence for Late Silurian–Early Devonian metamorphism correlative with the Klakas orogeny has not yet been recognized in northwestern Canada.

Geological data from the Saint Elias Mountains are consistent with a model for the Alexander terrane to have evolved in proximity to the northern Caledonian-Appalachian convergent margin. For example, distinctive Middle Devonian trilobite fossils in southwest Yukon are elsewhere only known in Givetian strata of the Canadian Arctic Islands (Dodds et al., 1993). Underlying the Icefield assemblage, Cambrian-Ordovician sandstones in the Saint Elias Mountains yield Neoproterozoic to Neoarchean detrital zircons with Hf isotopic compositions that strongly favor provenance connections with the Timanide orogenic belt and Fennoscandian Shield of northeast Baltica (Beranek et al., 2013). In the model of Beranek et al. (2013), the Alexander terrane formed part of an arc system that fringed the Uralian margin during the early Paleozoic. Cambrian-Ordovician rocks of the Eurasian High Arctic have Timanide and Fennoscandian provenance signatures similar to those of the Saint Elias Mountains and are also overlain by Old Red Sandstone-like successions (Gee et al., 2006; Lorenz et al., 2008a).

SAMPLES AND METHODOLOGY

Five rock samples of the Icefield assemblage in the Saint Elias Mountains were collected for detrital zircon studies. The suite includes three samples of quartzose sandstone from Yukon (09VL26, 09VL40, 09VL42) and one sample each of quartzose sandstone (09VLB27) and sandy limestone (09VLB28) from British Columbia (Fig. 1A).

Detrital zircon U-Pb geochronology was performed by laser ablation–inductively coupled plasma–mass spectrometry (LA-ICP-MS) at the Arizona Laserchron Center (ALC), University of Arizona (for analytical protocols, see Gehrels et al., 2008). Complete results and sample locations are given in the GSA Data Repository (Table DR1).¹ Detrital zircon U-Pb age results are presented in relative probability plots with stacked histograms (Fig. 2) made with the Isoplot 3.0 Excel macro of Ludwig (2003). Following data acquisition and reduction protocols at the ALC, analyses with >10% uncertainty or excessive discordance (>20% discordance or >5% reverse discordance) were rejected and not included in the probability plots.

Detrital zircon Hf isotope geochemistry was performed at the ALC using laser and ICP-MS routines similar to those used for U-Pb geochronology (for analytical protocols, see Cecil et al., 2011). In situ Hf isotope measurements were conducted on 176 dated zircons (16–47 grains analyzed from each sample; Table DR2 [see footnote 1]). Initial ¹⁷⁶Hf/¹⁷⁷Hf ratios are reported as $\varepsilon_{Hf(t)}$, which represents the isotopic composition at the time of crystallization relative to the chondritic uniform reservoir. The



Figure 2. Detrital zircon U-Pb probability density distribution-histogram plots of the Icefield assemblage. BC-British Columbia.

 $\epsilon_{\rm Hf(t)}$ values were calculated using the ^{176}Lu decay constant of Scherer et al. (2001) and the chondritic values of Bouvier et al. (2008). Hf model ages were approximated to the time of crystallization using $^{176}Lu/^{177}Hf = 0.015$ for the present-day crust (Goodge and Vervoort, 2006). Hf model ages provide an estimate for the timing of extraction of source rocks from a depleted mantle reservoir.

DETRITAL ZIRCON SIGNATURES

Icefield assemblage strata yield detrital zircons that range from 390 Ma to 3372 Ma (Figs. 2A–2E), with 405–490 Ma ages forming up to 92% of each sample. The youngest age groupings (three or more grains) in the three Yukon samples occur at 403–422 Ma, 417–440 Ma, and 424–432 Ma, respectively (Figs. 2A–2C), and with available fossil data from Dodds et al. (1993) suggest Upper Silurian to Lower Devonian stratigraphic ages.

¹GSA Data Repository Item 2013074, Table DR1, detrital zircon U-Pb age results; Table DR2, detrital zircon Hf isotope results, is available at www.geosociety.org /pubs/ft2013.htm, or on request from editing@geosociety.org, Documents Secretary, GSA, P.O. Box 9140, Boulder, CO 80301-9140, USA.

The two British Columbia samples each contain 390–395 Ma zircon age groupings (Figs. 2D–2E) and with the available fossil constraints suggest Middle Devonian stratigraphic ages.

Figure 3A shows the $\varepsilon_{\rm Hft0}$ values for 176 zircons that span 390–3372 Ma. As an age group, the 405–490 Ma detrital zircons can be divided into two subpopulations based on initial Hf isotopic composition and stratigraphic age. Upper Silurian to Lower Devonian strata contain 405–490 Ma zircons with mainly negative $\varepsilon_{\rm Hft0}$ values (84%) and Hf model ages that suggest Mesoproterozoic to Paleoproterozoic crust was involved in their genesis (Fig. 3B). Middle Devonian strata contain 405–490 Ma zircons with mostly positive $\varepsilon_{\rm Hft0}$ values (79%) and Neoproterozoic and younger Hf model ages (Fig. 3B).

DISCUSSION AND CONCLUSIONS

The detrital zircon Hf isotopic compositions of the Icefield assemblage record a major shift in the source of sediment between Late Silurian–Early Devonian and Middle Devonian time. This shift was likely the result of first-order changes in tectonic setting and paleogeography of the Alexander terrane. To explain these data, we interpret the Late Silurian–Early Devonian Klakas orogeny to manifest Scandian-Salinic deformation in the region of northeast Baltica, east of the Appalachian-Caledonian convergent margin (Fig. 4A). An unresolved problem concerning the Klakas orogeny is how the Late Cambrian–Early Silurian upper-plate setting for the Alexander terrane above a subduction zone fits in with the overall lower-plate setting for Baltica in the Scandian-Salinic collision. Arc magmatism in the Alexander terrane ceased by ca. 430 Ma (Gehrels and Saleeby, 1987), 15–20 m.y. prior to the Klakas orogeny, and the termination of subduction probably records a change in plate dynamics related to the closure of Iapetus, such as a subduction polarity reversal.

Upper Silurian to Lower Devonian molasse strata of the Alexander terrane are herein viewed as now-displaced equivalents of the Old Red Sandstone that covered much of the northern Appalachian-Caledonian orogen (cf. Soja and Krutikov, 2008). The negative $\epsilon_{_{Hf(t)}}$ values and Mesoproterozoic to Paleoproterozoic Hf model ages of 405-490 Ma detrital zircons in rock samples 09VL26, 09VL40, and 09VL42 accordingly support provenance connections with crustally contaminated granites in the Greenland, Svalbard, or British Caledonides. For example, both arc-type and anatectic granite suites in Greenland (GL in Fig. 4A) incorporated material from surrounding Precambrian basement rocks and yield high initial Sr isotopic ratios (0.709–0.711), magmatic zircons with $\epsilon_{\rm Hf(t)}$ values of –11.9 to –5.5, and Paleoproterozoic Hf model ages (Rehnström, 2010; Augland et al., 2012). Caledonian anatectic granites in Svalbard (SV in Fig. 4A) have peraluminous geochemical signatures, negative $\epsilon_{_{Nd(j)}}$ values, and inherited Proterozoic zircons that together provide strong evidence for crustal melting and contamination (Johansson et al., 2005). Although detrital zircon Hf isotopic data for Old Red Sandstone units are lacking, we note that our U-Pb age populations are broadly similar to relevant Silurian-Devonian molasse strata in Svalbard (Pettersson et al., 2010), Scotland (Phillips et al., 2009), and Severnaya Zemlya (Lorenz et al., 2008b).

The Devonian period was a time of rifting in northern Laurussia and transcurrent displacement within the Caledonides (Nikishin et al., 1996; Dewey and Strachan, 2003). Devonian rifts notably formed part of the extensive Slide Mountain–Angayucham backarc system, which accompanied an arc chain situated along much of the length of northwest Laurussia (Colpron and Nelson, 2011; Miller et al., 2011). We propose that this arc chain formed by a Devonian subduction polarity reversal along the northern extension of the Caledonian convergence system, following the terminal Scandian-Salinic collision that created Laurussia (Fig. 4B). The presence of Middle to Late Devonian arc-type rocks in southeastern Alaska (Gehrels and Berg, 1992; Saleeby, 2000; Cecil et al., 2011), locally possibly as old as Early Devonian (Nelson et al., 2012), along with existing biogeographic and paleomagnetic constraints, provide evidence for the Alexander terrane to comprise part of this magmatic system (Fig. 4B).

It seems likely that the opening of the Slide Mountain–Angayucham backarc rift led to the tectonic separation of the Alexander terrane from



Figure 3. (A) $\varepsilon_{Hf(t)}$ vs. U-Pb age diagram for all detrital zircons of the lcefield assemblage. (B) $\varepsilon_{Hf(t)}$ vs. U-Pb age diagram for Paleozoic detrital zircons of the lcefield assemblage. CHUR—chondritic uniform reservoir. The depleted mantle Hf evolution curves were calculated from values reported by Vervoort and Blichert-Toft (1999). Hf model ages or crustal evolution lines at 1.0, 2.0, and 3.0 Ga are plotted using ¹⁷⁶Lu/¹⁷⁷Hf = 0.015 (Goodge and Vervoort, 2006).



Figure 4. (A) Late Silurian–Early Devonian paleogeographic reconstruction based on the results of this study and Beranek et al. (2013). The Alexander terrane restores to a position to the east of the lapetan suture and within the Scandian deformation front of the Eurasian High Arctic. AA–Arctic Alaska, AX–Alexander terrane, CH–Chukotka, GL–Greenland, FW–Farewell terrane, SV–Svalbard, SZ–Severnaya Zemlya. (B) Middle to Late Devonian paleogeographic reconstruction modified from Colpron and Nelson (2011) and Miller et al. (2011). AG–Angayucham Ocean, SM–Slide Mountain Ocean. The Alexander terrane is positioned on the trenchward side of the Slide Mountain–Angayucham backarc rift system.

northern Laurussia. In addition to the generation of juvenile arc-type rocks, we predict that the timing of this separation is indicated by the relative absence of Caledonian-derived zircons in Middle Devonian sequences of the Alexander terrane (samples 09VLB27, 09VLB28). These samples instead contain 390–490 Ma detrital zircons with juvenile Hf isotopic compositions that are consistent with the erosion of local basement rocks. This hypothesis agrees with studies that consider the Alexander terrane to have evolved apart from the northern Laurussia cratons during the mid-Paleozoic to Mesozoic (Gehrels and Saleeby, 1987). Arc-type magmatism in southeastern Alaska ended by the Late Devonian, and the Alexander terrane evolved in a tectonically stable environment until the late Paleozoic.

IMPLICATIONS FOR FUTURE DETRITAL ZIRCON Hf ISOTOPE STUDIES

Building upon the conclusions of this investigation, we predict that future detrital zircon Hf isotope studies will be able to more accurately evaluate the crustal evolution and paleogeography of Cordilleran terranes. Paleozoic detrital zircon age populations that characterize the Icefield assemblage are also recognized in Silurian–Devonian strata of the Arctic Alaska-Chukotka terrane of Alaska (Amato et al., 2009), eastern Klamath terrane of California and Oregon (Grove et al., 2008), Chilliwack composite terrane of Washington and southern British Columbia (Brown et al., 2010), and Quesnellia of southern British Columbia (Lemieux et al., 2007). Along with Alexander, these Cordilleran terranes contain Devonian magmatic rocks and biogeographic signatures that suggest proximity to an arc system along the trenchward side of the Slide Mountain–Angayucham backarc rift. Detrital zircon Hf isotope studies of these terranes may shed new insights into their evolved or juvenile crustal sources and the overall significance of Cordilleran-Caledonian connections to global paleogeography.

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