# REE in carbonates from sediment-hosted Pb-Zn deposits, British Columbia, Canada

## Introduction

Paleozoic platform carbonate rocks of the Rocky Mountains host Mississippi Valley-type (MVT) deposits in the north, and in the south host MVT, magnesite barite, and Rare Earth Element (REE)-Ba-F deposits. Farther to the west of the Rocky Mountains, the platform carbonate rocks of the Kootenay arc pericratonic terrane also host MVT deposits in addition to vein- and replacement-type Zn-Pb  $(\pm Ag, \pm Au)$  deposits. Sparry dolomitization is commonly spatially associated with most of these deposits and is used as an exploration guide.

This is the first systematic study documenting REE variations in carbonate minerals from MVT deposits. In conjunction with tectonic, stratigraphic, structural, stable and radiogenic isotope, and petrographic studies it will help to constrain the genesis of MVT deposits and improve our knowledge of the metallogeny of the Kootenay arc and Rocky Mountains. It is based on more than 200 analyses by Laser Ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICP-MS), supported by electron microprobe work. Investigated samples are from the Munroe, Shag, Kicking Horse, Monarch, Coral and Robb Lake deposits in the Rocky Mountair belt, and from the Pend Oreille, Reeves MacDonald, Jersey-Emerald, HB, Duncan and Abbott-Wagner (a fracture-controlled replacement) deposits within the Kootenay arc.

In addition to high spatial resolution, LA-ICP-MS offers a further advantage over traditional whole rock analyses of carbonate concentrates: allowing the analyst identify compositional artefacts from inclusions and interstitial material through careful examination for sharp compositional changes within the time-resolved spectrum of each analysis. Thus, selection of carbonate grains free of visible mineral inclusions, followed by examination of corresponding LA-ICP-MS spectra permits identification and rejection of 'contaminated' analyses, retaining only the analyses representing the trace element signature of carbonates. The high spatia resolution of the LA-ICP-MS method requires a large number of individual analyses to ensure that results are representative.

## Geologic Setting

All of the deposits consider in this study are carbonatehosted. They are all MVT as characterized by Leach et al. (2005) and Paradis et al. (2007), with the exception of Abbott-Wagner, (fracturecontrolled replacement). All of Paleozoic plutons the occurrences studied are in British Columbia, except for Pend Oreille (Washington state, USA). Deposits selected for this study are in the Rocky Mountain belt and the Kootenay arc. Kicking Horse, Monarch and Shag are spatially associated with the Kicking Horse Rim, a faultbound Precambrian paleobathymetric high. Two deposits (Robb Lake and Coral) are located in the



Map showing the locations of the deposits studied in this work. The Robb Lake and Coral deposits are shown on the inset map.

northern Rocky Mountains, outside the area depicted in Figure 1 (see inset).

Biostratigraphic ages of sedimentary rocks hosting these deposits vary from lower Cambrian to Upper Devonian. Radiometric ages are unavailable for most of the deposits we have studied as it is notoriously difficult to obtain ages from low temperature, sediment-hosted deposits. Although most carbonate rocks hosting the deposits are recrystallized and deformed, the regional metamorphic grade is typically in the sub-greenschist or lower greenschist facies. These deposits do not contain calculate minerals except where they are located in the contact metamorphic aureoles of Mesozoic intrusions (e.g., Jersey-Emerald deposit).

## Analytical Methods

LA-ICP-MS analyses were carried out at the University of Victoria using a NewWave UP-213 (Nd-YAG laser, 213 nm) interfaced to either a Thermo-Fisher Xseries 2 quadrupole ICP-MS or an Agilent 8800 triple quadrupole ICP-MS. Repeated analyses of the calibration standards (NIST 615, NIST 613, NIST 611) showed no measurable drift in instrument sensitivity over the course of each analytical session. The content of elements in the BCR2g standard thus obtained a within 15% (relative) of the GeoReM preferred value (Jochum et al., 2005) and repeated measurements over the course of each analytical session showed that analytical precision for the REE+Y is within 10% relative.





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## Guide to Interpretation



(a) PAAS-normalized REE profiles for seawater from the western North Pacific (Alibo and Nozaki, 1999) showing the distinct negative Ce anomaly and LREE depletion relative to HREE, pore water from marine sediments on the Californ margin (Haley et al., 2004), hydrothermal yent fluids from Manus Basin (Craddo et al., 2010), sparry calcite from the Dirtlow open-pit in the South Pennine Orefie (Bau et al., 2003), cold seep carbonates from northern Apennines, and calcite fror the volcanogenic massive sulphide Bracemac-McLeod deposits (Genna et al 2014). (b) A binary plot showing how true Ce and La anomalies are distinguish using the relationship between PAAS-normalized Ce and Pr anomalies, after Bau and Dulski (1996). (c) The effect of progressive contamination of stromatolit analyses by shale on the PAAS-normalized REE profile (van Kranendonk et al.,

#### Summary

1. This study shows that calcite and dolomite from the studied carbonate-hosted Zn-Pb deposits display significant differences in REE content, in the overall shape of PAAS-normalized REE plots and in the intensity of Eu, Ce, and Pr anomalies. In the case of fine-grained carbonates, these variations are probably due, at least in part, to modification of the REE signature corresponding original carbonates by later diagenetic, dolomitizing hydrothermal, contact metamorphic and/or regional metamorphic

2. Approximately 85% of carbonate minerals analysed plot outside of the shallow seawater field on PAAS-normalized Ce to Pr anomaly diagrams, indicating post-depositional re-equilibration, crystallization from hydrothermal fluids or metamorphic overprint

3. Strong positive Eu anomalies, particularly if associated with convex upward patterns are interpreted to mark the influence of hot (>200°C) hydrothermal fluids and correspond to carbonates precipitated from, or recrystallized in the presence of ambient hydrothermal fluids.

4. To fully appreciate the significance of these observations, including trends on Ce to Pr (PAAS-normalized) anomaly diagrams, the results presented here will need to be interpreted in conjunction with: the corresponding detailed major element analyses; detailed paragenetic studies; C, O, and Sr isotopes and clumped C and O isotopes; radiogenic isotopes; and fluid inclusior results

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