





BC Geological Survey GEOSCIENCE MAP 2015-01





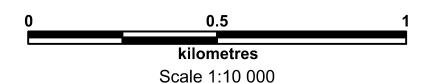
Geology, Geochronology and Ni-Cu-PGE Orebodies of the Giant Mascot Ultramafic Intrusion, Hope, Southwestern British Columbia

parts of NTS 092H/05 and 06

Sheet 1 of 2

M. J. Manor*, C. J. Wall*, R. M. Friedman*, J. Gabites*, G. T. Nixon**, J. S. Scoates*, and D. E. Ames***

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LEGEND

LATE CRETACEOUS

GIANT MASCOT ULTRAMAFIC SUITE (ca. 93 Ma)

Black and white mottled, medium- to coarse-grained hornblende gabbro with minor pyroxene (<15 vol %) and phlogopite (<5%); melanocratic varieties contain up to 75% stubby prismatic hornblende crystals.

Black to grey-brown, medium-grained to pegmatitic hornblendite and feldspathic hornblendite subhedral hornblende (90-98 vol %), minor pyroxene (<5%), rare phlogopite (<5%) and trace disseminated sulphide; forms a discontinuous pegmatitic zone up to 40m wide at the periphery of the ultramafic intrusion and also occurs in the pyroxenite.

Black to dark green, medium- to coarse-grained hornblende pyroxenite (websterite and orthopyroxenite), locally oikocrystic and feldspathic: anhedral-subhedral pyroxene (20-60 vol %) and hornblende (20-80%) with minor olivine (<5%) and trace sulphide; rare centimetre- to decimetre-scale modal layering.

pyroxenite), locally oikocrystic and feldspathic: subhedral pyroxene (55-90 vol %), olivine (5-35%), minor hornblende (<5%) and disseminated trace magnetite and sulphide. Dark brownish green, fine- to medium-grained, locally oikocrystic peridotite (Iherzolite and

Dark brownish green, medium-grained pyroxenite (websterite, olivine websterite and ortho-

harzburgite): subhedral olivine (40-80%), pyroxene (30-60%), minor hornblende (<10%), net-

textured sulphide (pyrrhotite, pentlandite and chalcopyrite, 2-25%) and trace chromian spinel. Dark green to yellow-brown, fine- to medium-grained, equigranular dunite: subhedral olivine (90-95%) and minor interstitial pyroxene (5-10%) with accessory chromian spinel and trace

SPUZZUM PLUTON (ca. 95 Ma)

Pale grey, medium-grained hornblende-biotite diorite and quartz diorite: subhedral plagioclase (An30-55; 40-60%), pyroxene (10-35%), biotite (5-20%), hornblende (5-15%) and accessory magnetite and ilmenite, and trace disseminated sulphides; entrains sparse ultramafic xenoliths and locally exhibits a moderate to strong foliation.

Pale grey to pinkish grey, medium-grained gabbronorite, diorite and minor norite: subhedral

Pale grey to pinkish grey, medium-grained, biotite-hornblende gabbronorite, diorite and minor norite: subhedral calcic plagioclase (An50-70; 45-75 vol %), pyroxene (15-35%), hornblende (5-20%), biotite (<5%), quartz (0-15%), Fe-Ti oxides (ilmenite and magnetite<2%) and trace disseminated sulphides; weak to moderate foliation.

calcic plagioclase (An50-70; 45-75 vol %), pyroxene (15-35%), quartz (0-15%), hornblende (<5%), Fe-Ti oxides (ilmenite and magnetite<2%) and sulphides (pyrrhotite, pentlandite, chalcopyrite<5%); weak to moderate foliation; rare centimetre-scale modal layering and subround xenoliths of pyroxenite and hornblende pyroxenite.

Garnetiferous gabbronorite-diorite adjacent to small raft of Settler Schist in the western part

UPPER TRIASSIC

SETTLER SCHIST

Dark grey, finely to coarsely crystalline, metasedimentary schist containing garnet, staurolite, kyanite, biotite, plagioclase and quartz, minor graphite and trace pyrite.

MAP SYMBOLS

Geological contact, defined

Geological contact, approximate Geological contact, inferred Modal layering, inclined Dyke (hornblendite or hornblende gabbro), inclined, vertical Foliation or schistosity, inclined Foliation or schistosity, inclined, vertical (Aho, 1954)

Foliation induced by shearing, inclined

U-Pb zircon date (Ma, 2 sigma; determined by R. M. Friedman) ⁴⁰Ar/ ³⁹Ar date (Ma; H, hornblende; L, large crystal; B, biotite; x, xenolith; inverse isochron dates determined by J. Gabites)

Foliation induced by shearing, inclined, vertical (Aho, 1954)

Orebody (projected to surface) Outcrop examined

Outcrop too small to show at map scale Flooded land (swamp)

Stream or river Contour (100m)

Contour (20m) Road (dirt)

Aho, A.E., 1954. Geology and ore deposits of the property of Pacific Nickel Mines near Hope, British Columbia. PhD Thesis, University of California, Santa Barbara,

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Christopher, P.A., and Robinson, J.W., 1975. Pride of Emory Mine (92H/SW-4). In: Exploration and Mining in British Columbia 1974, pp. 105–113.

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Colpron, M., and Nelson, J.L., 2011. A digital atlas of teranes for the Northern Cordillera. British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Geofile 2011-11.

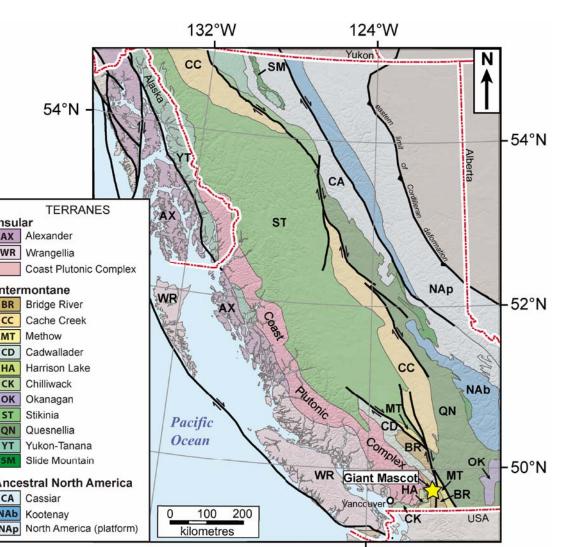
Manor, M.J., 2014. Convergent margin Ni-Cu-PGE deposits: geology, geochronology, and geochemistry of the Giant Mascot magmatic sulphide deposit, Hope, British Columbia. MSc Thesis, The University of British Columbia, Vancouver, British

Manor, M.J., Scoates, J.S., Nixon, G.T., and Ames, D.E., 2014. Platinum-group mineralogy of the Giant Mascot Ni-Cu-PGE deposit, Hope, BC. In: Geological Fieldwork 2013, British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Paper 2014-1, pp. 141-156.

McLeod, J.A., 1975. The Giant Mascot ultramafite and its related ores. MSc Thesis, The University of British Columbia, Vancouver, British Columbia, 123 p. Mitrovic, I., 2013. Evolution of the Coast Cascade Orogen by tectonic thickening, and

magmatic loading: the Cretaceous Breakenridge complex, southwestern British Columbia. MSc Thesis, Simon Fraser University, Vancouver, British Columbia,

Pinsent, R.H., 2002. Ni-Cu-PGE potential of the Giant Mascot and Cogburn ultramaficmafic bodies, Harrison-Hope area, southwestern British Columbia (092H). In: Geological Fieldwork 2001, British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Paper 2002-1, pp. 211-236.



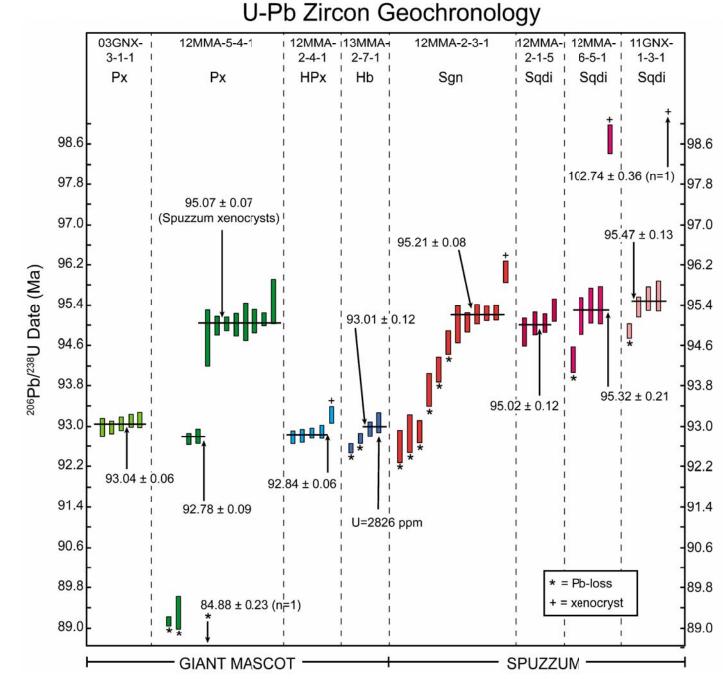
Tectonostratigraphic terranes and major faults of British Columbia and southeastern Alaska showing location of Giant Mascot (modified from

Giant Mascot Magmatic Ni-Cu-PGE Sulphide Deposit

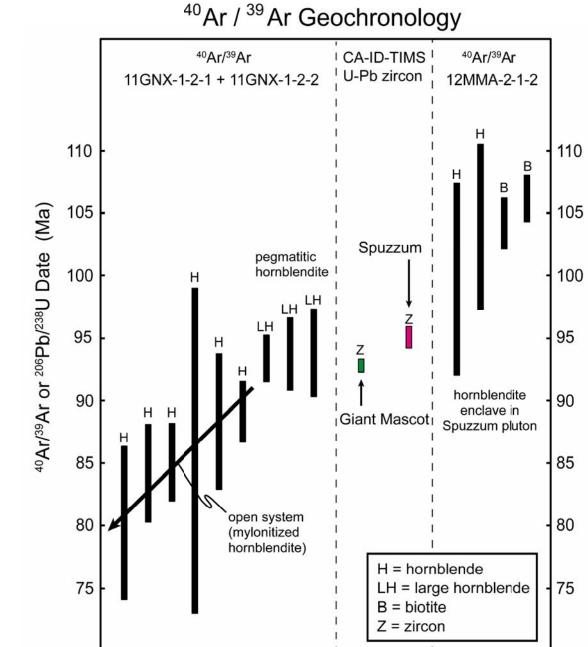
The Giant Mascot Ni-Cu-PGE deposit remains British Columbia's only past-producing nickel mine (1958-1974) with 4.2 million metric tonnes of ore grading 0.77 wt% Ni and 0.34 wt% Cu, minor Co, Ag and Au, and unreported platinum group elements (PGE; Christopher and Robinson, 1975). The deposit belongs to an increasingly important class of under-explored, orthomagmatic Ni-Cu±PGE sulphide deposits located at convergent margins. These arc-related, ultramafic-mafic intrusions, which are known to host economic mineralization, generally have small footprints (~200 m² to 10 km²) and may exhibit crudely concentric lithological zoning. Ultramafic cumulates and associated gabbros characteristically contain hornblende and orthopyroxene. The occurrence of cumulus orthopyroxene in ultramafic rocks distinguishes these intrusions from those of Alaskan-type affinity in the supra-subduction zone environment. The geology and geochronology of Giant Mascot summarized herein is taken from recent work by Manor (2014); the geological map showing the major rock units is modified from original mapping by Aho (1954).

The Giant Mascot ultramafic intrusion forms an elliptical plug (~3 x 1.3 km) at the southeastern margin of the Coast Plutonic Complex, the locus of Late Jurassic to Paleogene arc magmatism in the northern Cordillera. The intrusion is hosted by the Upper Triassic Settler schist and Late Cretaceous Spuzzum pluton. Regional metamorphism (mid- to lower crust) reached upper amphibolite grade at ~91-86 Ma (Mitrovic, 2013). The Giant Mascot ultramafic rocks are remarkably fresh and comprise dunite, hornblende-bearing peridotite (harzburgite, lherzolite) and pyroxenite (hornblende-bearing orthopyroxenite, websterite), hornblende pyroxenite and hornblendite with minor phlogopite and plagioclase, and accessory chromite, Fe-Ti oxides, apatite, zircon and rutile. Olivine-rich cumulates occupy the core of the crudely zoned intrusion and pegmatitic hornblendite-hornblende gabbro forms a discontinuous rim. The sulphide ores (mainly pyrrhotite, pentlandite and chalcopyrite) exhibit unambiguous magmatic textures and are predominantly hosted in dunite, peridotite and pyroxenite.

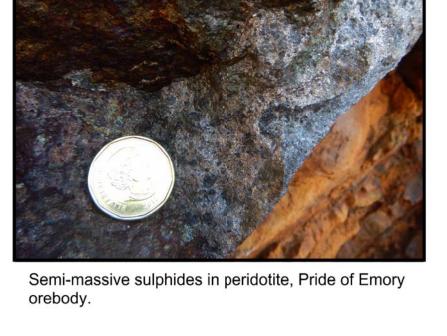
U-Pb zircon geochronology of pyroxenite and hornblendite has yielded a Late Cretaceous age (ca. 93 Ma) for crystallization of the Giant Mascot intrusion, and a slightly older but statistically distinct age (ca. 95 Ma) for the host Spuzzum diorite, in agreement with observed contact relationships. 40Ar/39Ar dating of a mylonite cutting hornblendite-hornblende gabbro at the rim of the ultramafic intrusion reveals open system loss of radiogenic argon starting shortly after consolidation (peak metamorphism ~91-86 Ma; Mitrovic, 2013). The new geochronological results indicate the timing of sulphide mineralization (ca. 93 Ma) and establish the Giant Mascot ores as the world's youngest known, orthomagmatic Ni-Cu-PGE deposit.



U-Pb zircon geochronology results for chemical abrasion ID-TIMS dating of Giant Mascot ultramafic rocks and Spuzzum diorite. Each bar is an analysis of an individual zircon grain showing extents of error (2 standard deviations). Black horizontal lines indicate weighted mean ²⁰⁶Pb/ ²³⁸U crystallization ages for each sample. Rock unit codes correspond to those on the geology map.

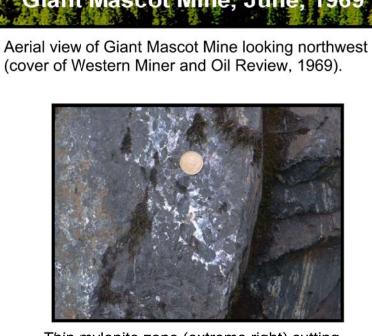


Comparison of U-Pb zircon [chemical abrasion ID-TIMS] and ⁴⁰Ar -³⁹ Ar (inverse isochron) dating results for Giant Mascot and Spuzzum intrusions. Each bar is an individual analysis showing extents of error (2 standard deviations). Hornblendite enclaves occur in the Spuzzum pluton.

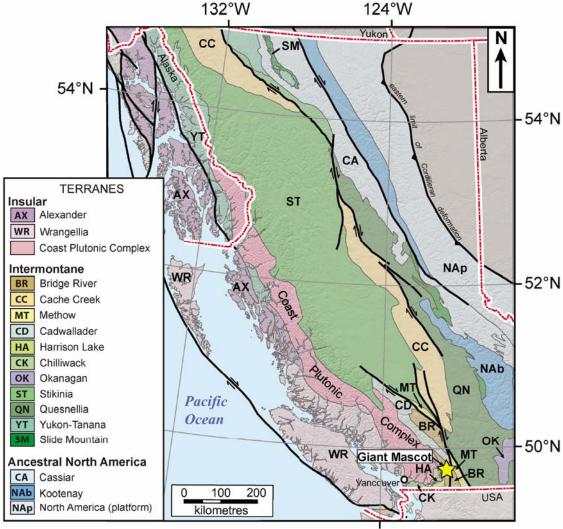




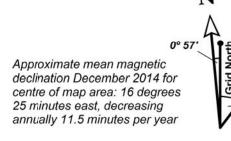
Mineralized pyroxenite exhibiting partially disrupted, weak modal layering. Hammer handle is 50 cm long.



Thin mylonite zone (extreme right) cutting pegmatitic hornblendite-hornblende gabbro.



Colpron and Nelson, 2011).



Base Map Information

Base map produced from digital TRIM (Terrain Resource Inventory Management) 1:20 000-scale topographic database, British Columbia Ministry of Forests, Lands and Natural Resource Operations. North American Datum 1983 (NAD83). Universal Transverse Mercator Projection (Zone 10). Elevation

in metres above mean sea level. Contour interval 20m.

Digital Cartography

G. T. Nixon M. J. Manor T. Barry

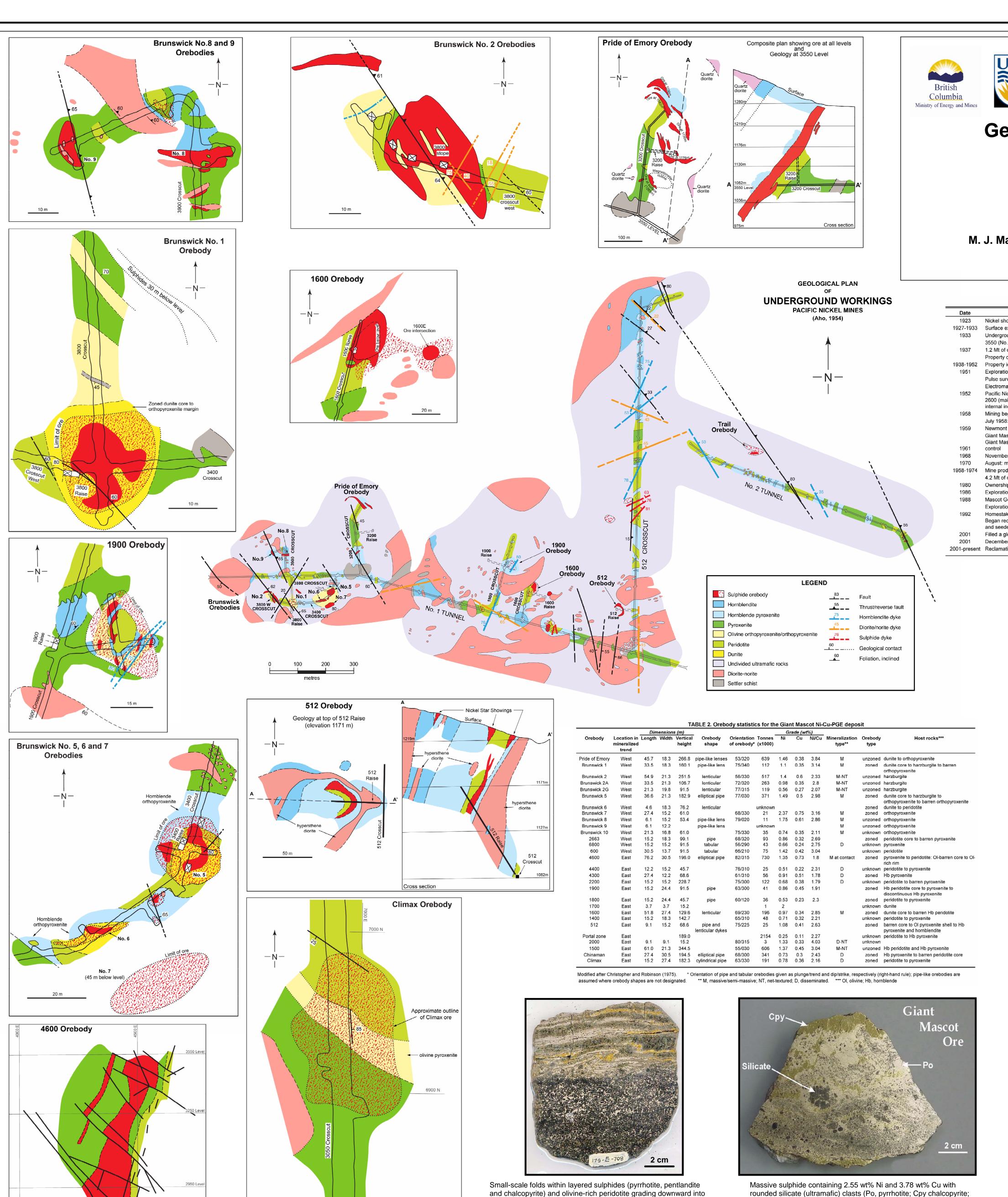
Recommended Citation

Manor, M. J., Wall, C. J., Friedman, R. M., Gabites, J., Nixon, G. T., Scoates, J. S., and Ames, D. E., 2015. Geology, geochronology and Ni-Cu-PGE orebodies of the Giant Mascot ultramafic intrusion, Hope, southwestern British Columbia, British Columbia Ministry of Energy and Mines, British Columbia Geological Survey Geoscience Map 2015-01, scale 1:10 000 (2 sheets).

Acknowledgments

We thank Wes Harmon, Alex Colyer and Lauren Harrison for help with mapping and sample collecting during 2012 and 2013 summer field seasons. The staff of Valley Helicopters in Hope are thanked for providing safe and efficient flights; and Barrick Gold and Fraser Valley Dirt Riders Association for access to their property. McKay Edwards kindly provided detailed weather reports at the mine prior to the 2013 field season. Helpful comments by Jon Scoates and Alex Zagorevski, Geological Survey of Canada, improved the map presentation. Funding for this project is provided by Natural Resources Canada Targeted Geoscience Initiative 4 (TGI-4) Program administered by the Geological Survey of Canada and a Society of Economic Geologists Canada Foundation Student Research Grant awarded to Matthew Manor in 2013.





dunite with net-textured to disseminated sulphides. Sample was

sample 179-E-708; Aho, 1956).

(McLeod, 1975)

collected at the contact between barren dunite and barren peridotite

where sulphide layers are steeply dipping (Pride of Emory orebody,

3550 Level adit area, sample RHP01-076; photo by R. H. Pinsent).



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TABLE 1. Giant Mascot mine and exploration history Nickel showings first discovered at Pride of Emory by Carl Zofka 1927-1933 Surface exploration B.C. Nickel Co. B.C. Nickel Mines Ltd. Underground development began 3550 (No. 1 tunnel) and 3275 (No. 2 or Chinaman tunnel) developed 1.2 Mt of ore mined at 1.38% Ni and 0.50% Cu Property closed down - all drill core lost Pacific Nickel Mines Ltd. 1938-1952 Property idle due to poor market conditions and World War II Exploration programs resumed Pulse survey conducted by Newmont Exploration Co. Electromagnetic and magnetometer surveys conducted by McPhar Geophysics Pacific Nickel Mines Ltd. and Newmont Mining Corp. of Canada merge Western Nickel Mines Ltd. 2600 (main haulage), 2950, and 3250 levels developed; levels connected with internal inclined shaft Mining began July 1958: Closed due to market conditions Newmont Mining Corp. sold property interest to Giant Mascot Mines Ltd. Giant Mascot Mines Ltd. opened mine as salvage operation Giant Mascot Mines Ltd. purchased Pacific Nickel Mines Ltd. and gained full Giant Mascot Mines Ltd. November: mine closed due to collapse of a Brunswick stope August: mine closed due to mill burning down 1958-1974 Mine production 4.2 Mt of ore mined at 0.77% Ni, 0.34% Cu, minor Co, Ag, Au Ownership of mine transferred to Mascot Gold Mines Ltd. Mascot Gold Mines Ltd. Exploration programs resumed with a focus on Au and PGE mineralization International Corona Corp. Mascot Gold Mines Ltd. was acquired by International Corona Corp. Exploration continued (results not released to the public) Homestake Canada Ltd. acquired International Corona Corp. Homestake Canada Ltd. Began reclamation of the Giant Mascot Mine, sealed portals, groomed dumps, and seeded tailings Filled a glory hole exposing the Pride of Emory and Brunswick deposits 2001-present Reclamation of the Giant Mascot Mine by Barrick Gold Corp.

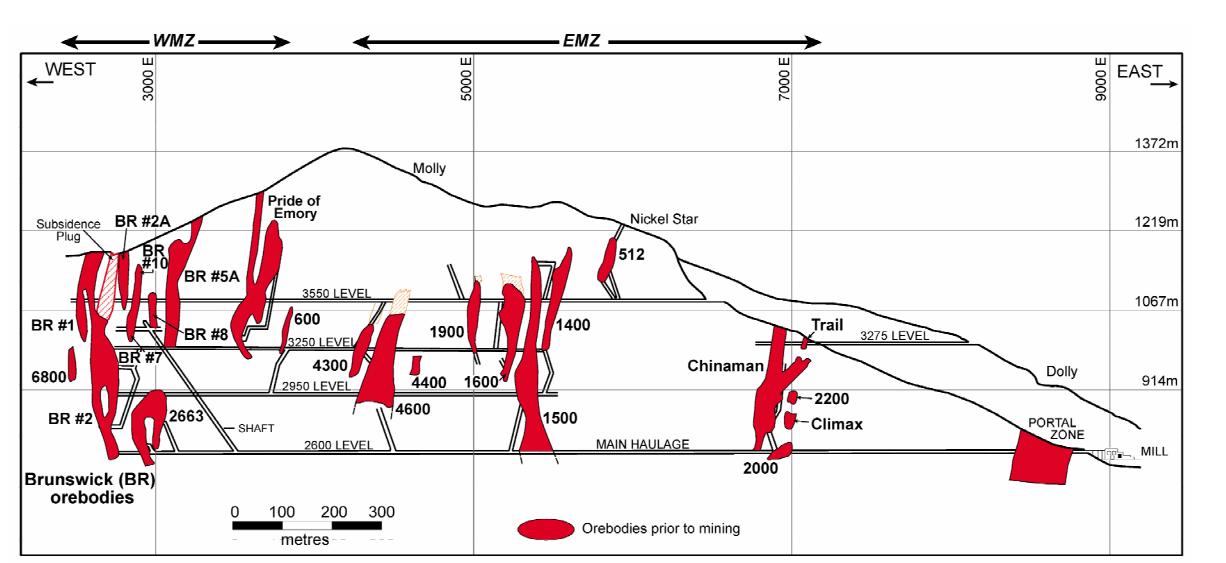
Giant Mascot Ni-Cu-PGE Orebodies

The Giant Mascot orebodies were discovered in 1923 by Carl Zofka at the Pride of Emory showing. Underground development and mining began in the 1930s and continued intermittently under various operators until mine closure in 1974 (Table 1; and Pinsent, 2002). Statistics for 28 orebodies are given in Table 2, and detailed plans and sections of selected orebodies are reproduced from Aho (1954, 1956), Clarke (1969), and McLeod (1975).

The main mineralized zone trends N75°W through Zofka Ridge. Sulphide ore zones have steep attitudes and form pipe-like, lensoid and tabular bodies measuring 4-76 m in length, 4-31 m in width and 15-345 m in vertical extent. The orebodies are classified as zoned or unzoned based primarily on ore textures. Zoned ore shoots are concentrically zoned, predominantly pipe-like bodies with a core of massive or semi-massive ore surrounded by a shell of disseminated to net-textured sulphide (e.g. Brunswick No. 1, 5 and 512). Massive and semi-massive mineralization containing up to 50 vol% crystals of olivine and/or orthopyroxene occurs in unzoned orebodies that are lensoid or pipe-like in shape (e.g., Brunswick No. 2, 8 and 9). The cores of both types of orebody are intimately associated with dunite, olivine-rich peridotite or less commonly orthopyroxenite, and locally contain rounded to irregular inclusions of ultramafic wallrocks and noritic schlieren. In addition to the textural variation of sulphides, zoned orebodies commonly exhibit an outward zonation from olivine-rich to olivine-poor host rocks. Unzoned massive orebodies may exhibit layered sulphides at their margins with rare, small-scale folds that formed while the cumulates were still mobile (e.g., Pride of Emory). Contacts between orebodies or their ultramafic host rocks and barren wallrocks are sharp to gradational over several centimetres, and are accompanied by distinct textural and mineralogical differences (Aho, 1956).

The highest Ni grades (~5 wt%) are associated with dunite-hosted ore (e.g., Brunswick No. 5 and 6; Aho, 1956). Overall, the sulphide ores have high tenors (3-14 wt% Ni, 0.1-17.1 wt% Cu, 84 ppb to 5 g/t total PGE; recalculated to 100 wt% sulphide). Iridium-group PGE contents distinguish two magma types in the western mineralized zone (WMZ) and eastern mineralized zone (EMZ). Platinum-group minerals are predominantly bismuthotellurides (moncheite, merenskyite, palladian melonite) and minor arsenide (sperrylite), derived via exsolution from base metal sulphides during cooling (Manor et al., 2014; Manor

The geometry, internal features and contact relationships of the ore-bearing structures are consistent with them representing magmatic conduits, a possibility originally recognized by Aho (1956). Certain orebodies are cored by virtually barren pyroxenite (e.g., 1900), potentially indicative of multiple injection; others have arcuate shapes indicating that they may have formed by wedging out of injected crystal-rich magma in blind conduits, or possibly through partial collapse and infilling of channelways by wallrock conditions, pose a significant challenge to exploration, yet the presence of PGE-enriched, high tenor sulphides underscores the potential for economic Ni-Cu-PGE deposits in convergent margin



Longitudinal west-east section of the Giant Mascot sulphide deposit distinguishing orebodies of the western mineralized zone (WMZ) from those of the eastern mineralized zone (EMZ). Section extends from the Brunswick (BR) orebodies in the west to the Portal Zone and mill in the east. Note the two topographic profiles in the east (modified after Clarke, 1969; Christopher and Robinson, 1975).

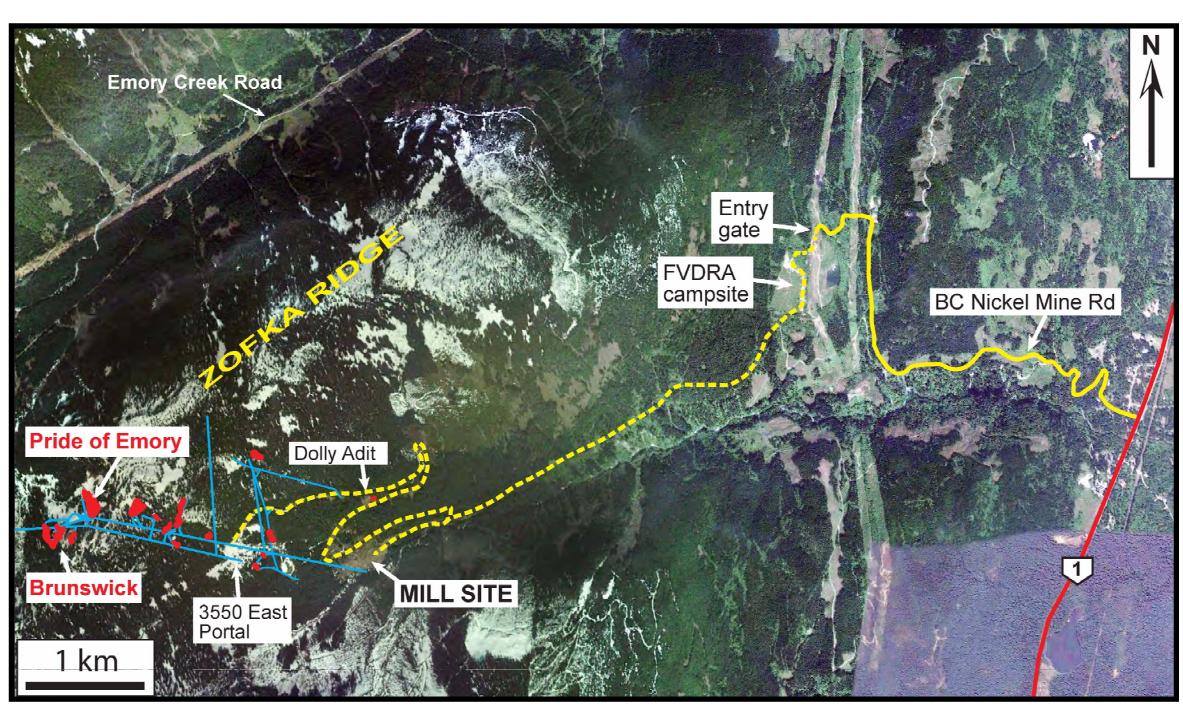


Image of Giant Mascot mine showing major orebodies and tunnels (projected to surface), and access roads (solid) and trails (dashed) to the mine site (FDRA, Fraser Valley Dirt Rider's Association; Google Earth, 2010; after Manor, 2014).