

#### INTRODUCTION

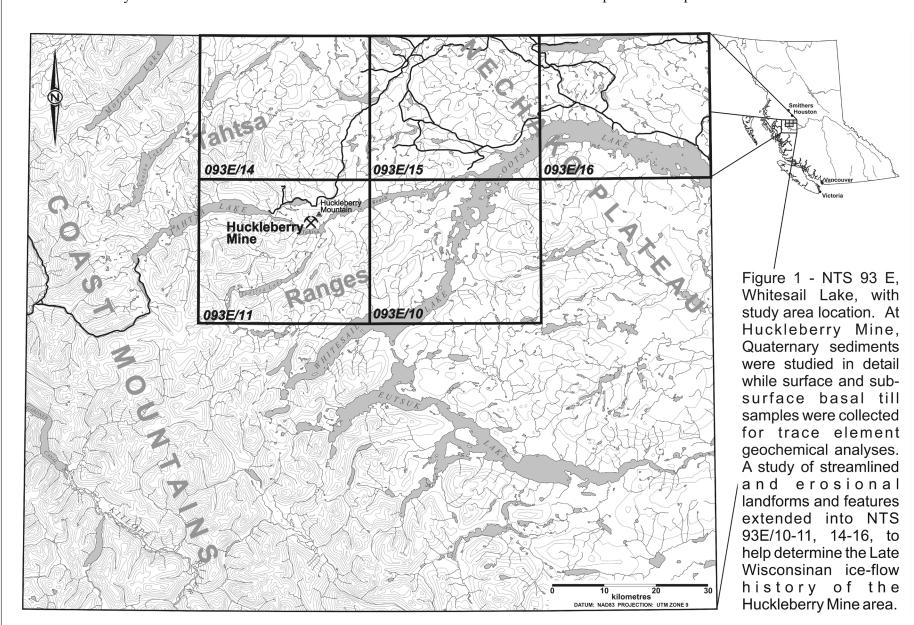
on the Huckleberry Mine property in northwest British Columbia km wide, oriented northwest-southeast, with peaks ranging from (Figure 1). Glacial history and ice flow studies were also completed 2100 to 2400 m in elevation. The Tahtsa Ranges are divided roughly on the property (north-central NTS 93E/11) and in the surrounding east-west by major valleys that are occupied by large lakes (e.g. area (NTS 93E/10-11, 93E/14-16), which included detailed Tahtsa, Troitsa, and Whitesail Lakes). These lakes are prominent stratigraphic and sedimentologic studies conducted at 12 exposures features of the Tahtsa Ranges and occur at relatively high elevations on the Huckleberry Mine property. Ferbey and Levson (2001a, b, c) (784 to 930 m) therefore reducing the overall relief of the area provide an overview of the 2000 field program and preliminary (Holland, 1976). Valley bottoms and mountain flanks are forested interpretations of the resulting till geochemical data sets and report on and have thick sequences of Quaternary sediments with little bedrock the complex ice-flow history of the Huckleberry Mine area.

This surficial geology map is a product of the 2000 field survey of the Huckleberry Mine area and interpretations of 1:15,000 scale aerial Huckleberry Mine (MINFILE No. 093E 037) is located on the north **PREVIOUS WORK** photographs. The terrain mapping conventions used here are based side of Tahtsa Reach, approximately 120 km southwest of Houston, The surficial geology of the study area was first mapped by the of the glacial and de-glacial phases of the Cordilleran Ice Sheet. on those of Howse and Kenk (1997), and focus primarily on surficial British Columbia, and is accessed via Forest Service and private mine materials and surface expression. A composite stratigraphic section roads (Figure 1). The mine site is located in a poorly drained, boggy, (Figure 2), based on the detailed stratigraphic and sedimentologic arcuate, U-shaped valley (averaging 1015 m elevation), adjacent to Huckleberry Mine property was produced by New Canamin sediments is provided in Figure 2 and in the map unit descriptions. studies conducted on the mine property, is included to assist in the southern flank of Huckleberry Mountain (1543 m). Huckleberry characterizing the Quaternary sediments of the Huckleberry Mine Mine is a porphyry copper-molybdenum open pit mine, with minor area.

LOCATION AND PHYSICAL SETTING

the Nechako Plateau to the east and the Coast Mountains to the west During the 2000 field season, geochemical studies were conducted (Figure 1). This area, the Tahtsa Ranges, is a belt of mountains 16-24 exposure. Upper slopes and peaks extend into sub-alpine and alpine environments

recoverable amounts of gold and silver, and has a production rate of 21,000 tonnes/day ore (Huckleberry, 2000). A more detailed discussion on local bedrock types and the style of mineralization can The Huckleberry Mine area falls within the transition zone between be found in the map unit descriptions. Other mineral occurrences



MINFILE N	UMBER	
093E	037	
093E	038	
093E	039	
093E	086	
093E	113	١

shown on the map.

and streamlined landforms are not included.

events.

to Holocene in age.

QUATERNARY GEOLOGY

NAME	STATUS	COMMODITY	DEPOSIT TYPE	
HUCKLEBERRY	Producer	Cu, Mo, Ag, Au	Porphyry Cu ± Mo ± Au	
LEN 45	Showing	Cu	Porphyry Cu ± Mo ± Au	
LEN 56	Showing	Cu, Zn	Polymetallic veins Ag-Pb-Zn±Au	
WEE	Showing	Cu, Mo	Porphyry Cu ± Mo ± Au	
HITING CREEK (SWEENEY)	Showing	Cu, Mo	Porphyry Cu ± Mo ± Au	

Resources (1993). In both cases, these maps are limited to the delineation and interpretation of surficial sediments only; *Ice-flow History* 

geology and glacial and geomorphic features of NTS 93 L and M. event followed by the east to northeast event. Again to the north, a detailed report on the Quaternary geology and

have produced radiocarbon ages that range from approximately subsequent events. 27,000 to >45,000 years B.P. (Levson and Giles, 1997; Levson et al.,

Although the Quaternary landscape depicted on this map is

found near Huckleberry Mine are listed in Table 1, with locations colluvial activity, peat forming processes and soil development. In comparison with areas mapped adjacent to it, a notable characteristic of the Huckleberry Mine area is an absence of glaciofluvial and glaciolacustrine deposits at the surface. The following is a discussion British Columbia Ministry of Environment, Lands and Parks (1976) pertaining to the Huckleberry Mine area, and the resulting general at a scale of 1:50 000, while a more detailed 1:10 000 scale map of the sediment types deposited. A more detailed description of these

identification and interpretation of ice flow indicators such as striae The Huckleberry Mine area has a complex ice-flow history. The orientation of streamlined landforms initially suggests a dominant ice flow direction of east to northeast during the last glacial maximum. Other Quaternary geological studies have been conducted in areas However, field investigations of these landforms and small scale adjacent to the Whitesail Lake map area (NTS 93 E). Immediately features on bedrock weakly exposed within them suggest a more north, Tipper (1994) and Levson *et al.* (1998) discuss the Quaternary complex ice flow history with an earlier west to southwesterly flow

till geochemistry of the Babine Copper Belt (NTS 93L/9, 16; M/1, 2, Based on the relative degree of landform preservation, the timing of 7-8) is provided by Levson (2001) and Levson *et al.* (2002). Similar opposing ice flow events inferred from cross-cutting and/or studies by Levson and Giles (1997), Levson *et al.* (1999), Plouffe and superimposition relationships, and the absence of all but west to Levson (2001), Levson and Mate (2002), focus on areas bordering to southwest ice-flow indicators at high elevation sites (>1500 m), the the northeast, east, and southeast, within NTS 93 K, F, and C west to southwest event appears to be earlier than the east to northeast respectively. Stumpf *et al.* (2000) discuss the timing of the Fraser event, and the product of a thicker ice sheet. This west to southwest Glaciation in central British Columbia, and associated ice-flow event is therefore interpreted as the dominant ice-flow direction during the Fraser Glaciation maximum which was followed in some low elevation areas by a weaker, possibly shorter lived, east to Quaternary geological studies within NTS 93 E have not yet revealed northeast ice flow event. Evidence of easterly ice flow out of the any Middle Wisconsinan (Olympia nonglacial) ages. However 100 Coast Mountains, which undoubtedly dominated during the km east of the Huckleberry Mine area, on Cheslaslie Arm, organic relatively long build-up phases of the Fraser Glaciation, is preserved rich sediments that underlie Late Wisconsinan (Fraser Glaciation) till at only a few sites as most evidence was likely obliterated by the

1998). As well, nonglacial sediments under till that have yielded These results are in general agreement with those discussed by finite radiocarbon ages occur in a number of other localities farther Levson *et al.* (1998, 1999), and Stumpf *et al.* (2000), and indicate the east in the Nechako Plateau (Plouffe and Jetté, 1997; Plouffe and presence of an ice dome in central British Columbia during the Fraser Levson, 2001) and approximately 200 km north of the study area at Glaciation maximum. With the development of this ice dome, ice the Bell Mine on Babine Lake (Harington et al., 1974; Levson, 2001). that once flowed east from the Coast Mountains during early Late The Quaternary sediments of the Huckleberry Mine area are believed Wisconsinan time now flowed west to southwest back towards and to be correlative with those found in areas adjacent to NTS 93 E over these peaks producing the observed west to southwest ice flow which overlie the dated organic rich and lacustrine deposits and are indicators. Towards the end of the Late Wisconsinan this divide therefore interpreted as being Late Wisconsinan (Fraser Glaciation) locally shifted west, back towards the Coast Mountains. As the ice sheet began to thin, ice flow again was influenced by topography producing the east/northeast ice flow indicators observed on air photos and in the field.

dominated by Late Wisconsinan morainal deposits, it has been For a more thorough discussion and presentation of ice flow data for modified significantly by post-glacial (or Holocene) fluvial and the Huckleberry Mine area, the reader is directed to Ferbey and

## Levson (2001c).

Advance Phase and Glacial Deposits Advance phase sediments in the Huckleberry Mine area are only seen At Stratigraphic Section 2000-1, multiple till units were identified. in vertical exposures and are composed primarily of clast supported, The units are distinguished primarily by differences in physical matrix filled, trough cross-bedded, large pebble to cobble gravels; characteristics such as clast size, clast content, matrix texture and bottom contacts are not exposed (e.g. Stratigraphic Section 2000-11). colour. The contacts between units are typically gradational over 20 These deposits, interpreted as advance phase glaciofluvial sediments, to 30 cm. The presence of multiple till units (each 2 to 6 m in Table 1 - Mineral Occurrences located near Huckleberry Mine. Locations, with MINFILE number, are shown on the map with a star. are observed peripheral to, and slightly down-elevation from, the thickness) could be attributed to: the complex ice-flow history of the central part of the arcuate valley that hosts the Huckleberry Mine area and changes in ice flow direction which resulted in a change in deposit. Up-section these deposits are interbedded with matrix provenance for individual till units; a change in depositional facies supported, massive to stratified, sandy-silt to silty-sand diamictons (e.g. from basal lodgement to basal meltout); or a combination of interpreted to be debris flow deposits.

> In the vicinity of the Main Zone Pit and in the East Zone Pit area, seen **Retreat Phase and Holocene Deposits** i vertical exposures in the valley bottom (approximately 1040 m), De-glacial phase deposits consist primarily of interbedded gravels are laminated silts and clays, pebbly muds, and massive fine sands. and sands. Gravel beds are clast-supported and include both These fine sediments are locally deformed (e.g. millimetre scale moderately to poorly sorted, matrix-filled, large pebble to cobble shears, normal and thrust faults), and contain small to medium gravels and open work, small to medium pebble gravels. Sands are pebble-sized dropstones (Stratigraphic Section 2000-2). They infill poorly sorted and mainly medium to coarse grained. These sands and a bedrock depression that opens to the east where a sediment dam gravels were observed on the Huckleberry Mine property in vertical nay have prevented free drainage to the east. Similar sediments are exposures only (e.g. Stratigraphic Sections 2000-5 to 7, and 2000observed in gully exposures on the southern flank of Huckleberry 10), directly over till. They are interpreted as glacio-fluvial Mountain (Stratigraphic Section 2000-3 and 4), at approximately sediments, deposited as ice retreated from the area. 1120 m. These fine grained sediments are interpreted as advance phase ice-marginal or sub-glacial glaciolacustrine sediments. At all Organic deposits are the dominate Holocene surficial material in the exposures, these glaciolacustrine sediments are interbedded with Huckleberry Mine area. They probably accumulated over much of massive, matrix supported diamictons. These poorly sorted the Holocene as indicated by wood fragments from a peat horizon sediments typically exhibit some degree of jointing and fissility, are overlying fan gravels that yielded a radiocarbon age of 7120 years moderately dense, can have striated clasts, and in some cases contain B.P. (±60 years B.P.) (Beta 158482). Organic veneers (**Ov**), swampfine silt/clay laminae or lenses up to 10 cm thick. These diamicton like in character, commonly border bodies of standing water, bogs, interbeds are interpreted as subaqueous glacigenic debris flow or and streams, while organic blankets (**Ob**), or bogs, are aerially more basal rain-out till deposits.

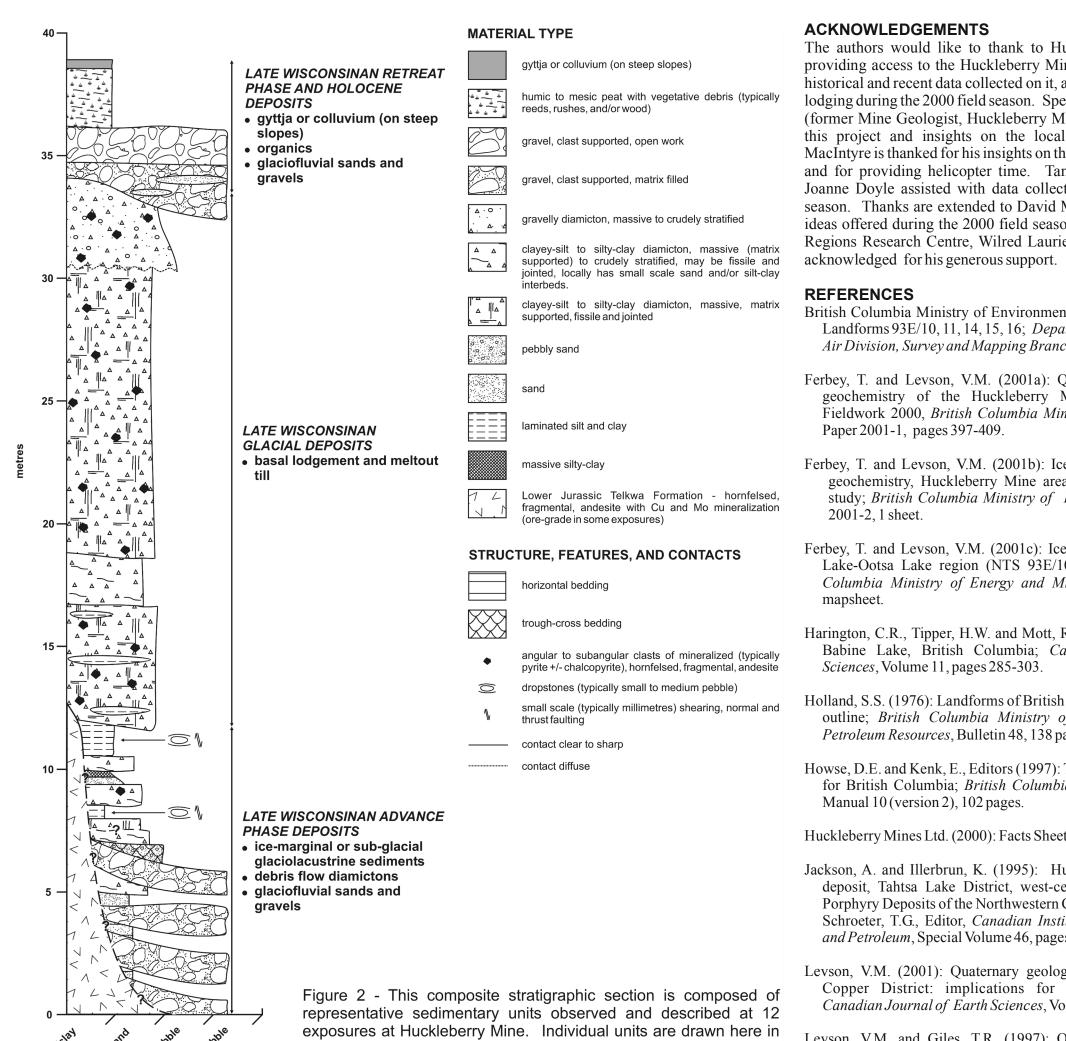
> biguitous cover in the Huckleberry Mine area. A product of Whiting Creek in the northeast corner of the map area and by Rhine comminution and reworking of bedrock and existing Quaternary Creek in the southwest. Both creeks drain aerially extensive, high deposits by glacial ice during the Fraser Glaciation, these deposits elevation areas to the north where glaciers continue to supply consist primarily of a massive and matrix supported, meltwater, helping account for these sizeable deposits. overconsolidated, sandy-silt diamicton, that is typically fissile and well jointed. These deposits are interpreted as basal tills and form Throughout the Holocene the main channel of Whiting Creek has thick blankets (Mb) in valley bottoms and on the lower and mid- moved laterally across the more open area at the mouth of the slopes of valley sides. Locally there are examples of gravelly-sand confined Whiting Creek valley. The decrease in stream competency diamictons, with undulating sand and gravel interbeds and as Whiting Creek exits the confined valley has resulted in sediment liscontinuous sandy and gravelly lenses. These sediments are aggradation and formation of a large fluvial fan (**Ff**; northeast corner nterpreted as basal meltout tills. In the vicinity of the East Zone Pit, of the map area). At some point during early Holocene time, Whiting basal tills reach a maximum known thickness of approximately 30 m Creek flow shifted across this developing fluvial fan, and flowed where they overlie striated, ore-grade bedrock.

> covered by organic (**Ob**) or colluvial blankets (**Cb**). Morainal dam at the west end of Sweeney Lake was topped and incised by the blankets typically overlie either bedrock (R) (e.g. Stratigraphic present day outlet. Seen in the field are poorly exposed, fine grained Sections 2000-1, 2000-3, and 2000-7) or glacial-fluvial gravels (e.g. sediments, in the vicinity of the Mount Sweeney and Sweeney Lake Stratigraphic Section 2000-11). The exception to this would be east Road intersection; these may be lacustrine sediments, perhaps also a of the Main Zone Pit and in the East Zone Pit area, where in addition product of increased lake levels. The Whiting Creek main channel to overlying bedrock, basal tills are found overlying the previously has since shifted back across the large fluvial fan towards the described glaciolacustrine sediments (Stratigraphic Section 2000-2). southeast and at present empties into Tahtsa Reach. At higher elevations, and on steeper slopes, till forms a thin veneer

(Mv) that mantles bedrock (R) or is found as a discontinuous cover in association with colluvial veneers (Cv).

extensive and substantially thicker, occupying large depressions in poorly drained areas. Other Holocene deposits in the area include Ill glacial deposits are the dominant Quaternary deposit and form a fluvial fans (Ff) and fluvial plains (Fp), such as those deposited by

southwest into Sweeney Lake. Deposition of the fluvial plain (**Fp**) on the east shore of Sweeney Lake may be the product of resultant Morainal deposits commonly occur at the surface but locally are increased lake levels which were likely maintained until a sediment



maximum true thickness.





Geological Survey Branch **OPEN FILE 2003-2** 

SURFICIAL GEOLOGY OF THE HUCKLEBERRY MINE AREA

- **NORTH-CENTRAL NTS 93E/11**
- T. Ferbey and V.M. Levson (PGeo) Cartography by T. Ferbey

SCALE 1:15 000

datum: NAD 83 projection: UTM Zone 9 elevation in metres above mean sea level; contour inteval 20 m Field survey carried out July to September, 2000

# GEOLOGY

**QUATERNARY SEDIMENTS** 

	osits: poorly sorted, angular gravels and sandy diamicts, found on moderate to steep slopes, associated with the down-slope transport and gravity processes) of weathered bedrock and pre-existing Quaternary sediments.
Cb	<b>Colluvial Blanket:</b> colluvial deposit >1 m in thickness, generally conforming to underlying topography but masking small scale irregularities; found mainly on moderately steep ( $\sim$ 15° to 20°), lower elevation slopes capping glacial units; locally forms talus deposits below steep slopes.
Cv	<b>Colluvial Veneer:</b> colluvial deposit 1 m or less in thickness, that conforms to minor surface irregularities of the underlying material; commonly found at higher elevations on steep to very steep slopes (~25° to 35°), where it is interspersed with areas of bedrock or morainal deposits.
iluvial Depos	its: silt, sand, and gravel deposited by streams and creeks; usually stratified, and with the exception of alluvial fans, is moderately to well sorted.
Fp	Fluvial Plain: thick fluvial deposit (>1 m) found on gently sloping surfaces; includes overbank deposits, terraces, abandoned and active channels of Whiting and Rhine Creeks, and the Sweeney Lake outlet; locally includes small areas of organic deposits along Whiting Creek; includes a deltaic deposit located at the east end of Sweeney Lake.
Ff	Alluvial Fan: thick deposits (>1 m) found at the mouth of a gully or confined valley; these deposits are dominantly fluvial in origin but likely include debris flow sediments.
)rganic Depo	sits: bog or swamp like, unforested, composed primarily of moss and organic soils, with a shallow water table found at or near the surface.
Ob	<b>Organic Blanket:</b> organic deposit >1 m in thickness, found in poorly drained areas that overlie thick morainal and fluvial deposits; typically fibric to mesic peats with identifiable organic material ranging in size from twigs and cones to tree limbs and stumps; can be capped by organic muds or gyttja; small ponds are present with small, low discharge, streams running through or from them; these deposits occupy relatively large areas east of Sweeney Lake and on the Huckleberry Mine property between the Main and East Zone Pits.
Ov	<b>Organic Veneer:</b> organic deposit 1 m or less in thickness; commonly borders ponds, lakes, and streams, but is also associated with bog (i.e. <b>Ob</b> ) margins; <b>Ov</b> is cross-hatched where other surficial units are also extensive.

### LATE PLEISTOCENE DEPOSITS

Morainal Deposits: typically massive, matrix-supported, sandy-silt, light brown to light grey, poorly sorted diamicts, originally deposited by or from glacial ice; these deposits are predominantly overconsolidated basal tills with vertical jointing and good sub-horizontal fissility; striated clasts are common; locally these tills can have a more sand rich matrix and can contain discontinuous lenses, and in some cases interbeds, of poorly sorted gravelly sands; at Huckleberry Mine, mineralized andesite and granodiorite clasts are common immediately adjacent to, and west of, the Main and East Zone pits; pyrite and chalcopyrite grains are also common here within the diamict matrix which is locally dark grey and clay-rich.

Morainal Blanket: morainal deposit >1 m in thickness, with a known maximum thickness of approximately 30 m in the vicinity of the Main and East Zone pits; generally conforms to underlying topography but masks small scale irregularities; these deposits form an extensive cover in valley bottoms and on the lower and mid-slopes of highland areas; gullied morainal deposits (**Mb-V**) have the same general characteristics but have been deeply incised by fluvial processes, and are typically found on the steeper flanks of Huckleberry Mountain and on the northern slopes of Tahtsa

Morainal Veneer: morainal deposit 1 m or less in thickness, that conforms to minor surface irregularities of the underlying material; typically a basal till that is found in higher relief areas (e.g. north shore of Tahtsa Reach and south of mine area) and higher elevations of Huckleberry Mountain; associated with bedrock outcrop, but can also contain small areas of colluvial and morainal blanket material; Mv is cross-hatched where other map units are also extensive

### **PRE-QUATERNARY BEDROCK**

Bedrock: large areas of bedrock outcrop are limited to the higher elevations of Huckleberry Mountain; flooding of the Nechako Reservoir, and resulting wave-wash processes, have created discontinuous bedrock outcrop along the shores of Tahtsa Reach: small areas of bedrock outcrop are also found locally in steeper colluviated areas above the north shore of Tahtsa Reach; bedrock of the Huckleberry Mine area is typically hornfelsed, fragmental, andesite of the Early Jurassic Telkwa Formation (MacIntyre, 1985); at Huckleberry Mine, mineralization occurs within a stockwork system of fractures and veinlets associated with two porphyritic hornblende-biotite-feldspar granodiorite intrusions (Late Cretaceous Bulkley Intrusions); these stocks are adjacent to or within the Main and East Zone pits (MacIntyre, 1985; Jackson and Illerbrun, 1995); although there is copper and molybdenum mineralization within both stocks, nearly all ore mined is from the adjacent volcanic rocks (S. Blower, pers. comm.)

NOTE: Composite delimiters on the map separate more extensive (single back-slash, /) and considerably more extensive (double back-slash, //) surficial units from less extensive units (e.g. Mv//R - Mv is considerably more extensive than R).

## SYMBOLS

# **GLACIAL FEATURES**

Irface till sample	•	roche moutonnee, crag and tail, flute (uni-directional)	ø
ub-surface till sample (collected using Becker-hammer drill)		ridge, flute (bi-directional)	ø
ratigraphic section site (section number)	2000-10	groove	/
ineral occurrence (MINFILE number)	093E 037	rat-tail	メ
eologic contact (approximate)		striae (uni-directional)	P
ctive pit outline	· · · · ·	striae (bi-directional)	P
ctive mine area			
ad			

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