

British Columbia Geological Survey Geological Fieldwork 1988

GEOLOGY OF THE CANTIN CREEK AREA QUESNEL RIVER (93B/16)

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KEYWORDS: Regional geology, Quesnel belt, Cantin Creek, stratigraphy, lithogeochemistry, alteration. gold.

INTRODUCTION

The Cantin Creek map area lies in the south-central part of the Quesnel belt, about 33 kilometres south of Quesnel and 55 kilometres northwest of Likely (Figure 1-20-1). It covers an area of approximately 100 square kilometres between latitudes 52°53' and 53°00' north and longitudes 122°22' to 122°08' west. The area was studied as part of a 4-year mapping project, the goal of which is to interpret the geological setting and evaluate the gold and copper resource potential along the central volcanic-intrusive axis of the Quesnel belt, (Bloodgood, 1987, 1988; Bailey, 1988; Panteleyev, 1988).

The Quesnel belt, previously known as the Quesnel trough, consists of Upper Triassic and Lower Jurassic basic to intermediate volcanic and volcaniclastic rocks, as well as coeval alkalic intrusions. The belt is bounded to the east by the Precambrian to Lower Paleozoic Snowshoe Group and to the west by the Permo-Carboniferous Cache Creek Group (Figure 1-20-2).

PREVIOUS WORK

The first geological investigation of the Quesnel belt dates back to 1887 when G.M. Dawson recognized Triassic volcanic rocks near Kamloops. Extensive regional mapping and local, detailed research were carried out only after the 1940s. In the 1970s, Fox (1975), Lefebure, Morton, Barr *et al.*, Hodgson *et al.*, (all 1976), Bailey (1978) and Preto (1979) described the alkaline nature of the plutonic and volcanic rocks of the region.

Early mining activity in the area was limited to placer gold operations. From the late 1960s, exploration for porphyry copper and copper-gold deposits and, more recently, mesothermal and epithermal gold-bearing systems has occurred, The discovery of several deposits, including the QR gold deposit (Fox *et al.*, 1986), is a direct result of these efforts.

LITHOLOGY

Due to the general sparseness of outcrop, fault offsets, and the limited size of the map area, correlation of map units is difficult. Fortunately there are two seemingly continuous horizons of volcanic wackes and one horizon of maroon basalt flows that may be used as markers. The sequence established here (Figure 1-20-3) is compatible with those of previous workers in the Quesnel belt. In this study Unit 1 is equivalent to Unit 1 of Bailey (1988) and Panteleyev (1988). Similarly Units 2 to 5 correspond to Bailey's Units 2A to 2H, and Unit 6 is part of his Unit 3. The instrusive Units 7, 8 ar d 9 are similar to those described by Bailey and Panteleyev, except that Unit 7 in the Cantin Creek area contains alkalic, mafic cumulate material as well as diorite and monzonite.

Unit 1 – Argillite: Dark grey to black, thinly bedded, locally with thin layers of fine-grained, pink to pale grey feldspathic wacke. The unit is exposed in the northwest part of the map area and as xenoliths within an intrusion of megacrystic quartzose syenite porphyry (Unit 9). The stratigraphic thickness is unknown due to the intrusion of the porphyry.

Unit 2 – Basalt Flows: Dark green, porphyritic with phenocrysts 2 millimetres in average diameter consisting of 3 per cent feldspar, 8 per cent pyroxene and minor olivine. The matrix is altered and contains carbonate and chlorite. The base of the unit is not exposed and the thickness is not known.

Unit 3 – Pyroxene-bearing Wacke: Maroon, coarsegrained; subrounded grains; locally well bedded otherwise massive, consisting of 15 per cent feldspar, 15 per cent

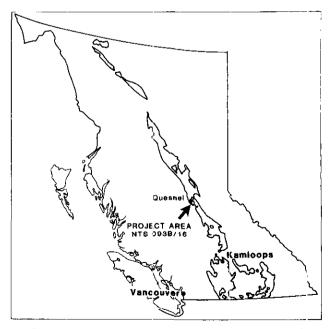


Figure 1-20-1. Location of Cantin Creek map-area in Quesnel terrane.

British Columbia Ministry of Energy, Mines and Petroleum Resources, Geological Fieldwork, 1988, Paper 1989-1.

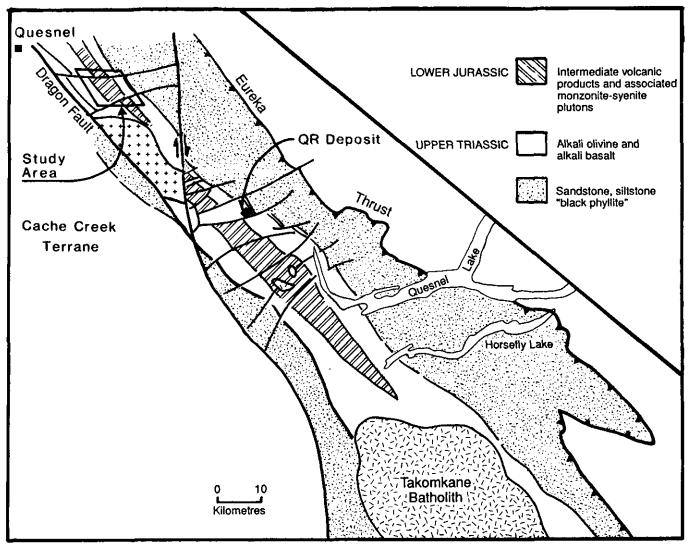


Figure 1-20-2. Regional setting of Cantin Creek map-area in Quesnel belt, (93A, B) (after Bailey, 1988).

pyroxene, 3 per cent iron oxides and minor amounts of lithic fragments. Towards the top of the unit the grain size decreases significantly and crossbedding features are more evident, the feldspar proportion increases and the unit assumes a pale greenish tinge. The thickness of the unit ranges from 80 to 270 metres.

Unit 4 – Maroon Basalt: Porphyritic with 45 per cent pyroxene, 12 per cent feldspar and 3 per cent olivine phenocrysts. At the base of the unit, sphene phenocrysts and ovoid amygdules of analcite and calcite are well developed. The unit is relatively continuous and varies in thickness from 770 metres in the southeast to 400 metres in the middle of the map area.

Unit 5 – Feldspathic Wacke: Maroon, consisting of subrounded to angular fine-grained fragments of feldspar, minor lithic fragments and iron oxides. The unit is capped by a thin layer of limestone. It is thickest in the southeast at 480 metres and thins towards the northwest.

Unit 6 – Polylithic Breccia and Feldspathic Tuffs: Breccias with feldspathic and heterolithic clasts from underlying tuffaceous rocks are dark green to maroon, consisting of 70 to 80 per cent feldspar, 10 to 20 per cent pyroxene, minor olivine and other minerals. In the southeast of the map area, breccias are most common. The thickness of the unit is not well defined. The section is 700 to 1300 metres thick but probably has some structural thickening.

Unit 7 – Pyroxenite, Gabbro, Diorite, Monzonite and Minor Syenite: This unit intrudes Unit 6. The mafic rocks are green due to extensive chloritization. Where sampled they contain 50 per cent phenocrysts consisting of 35 per cent pyroxene, 10 per cent feldspar, and 5 per cent olivine. The matrix composition is optically indeterminable due to alteration. The map unit is poorly exposed and is mainly defined by diamond drilling.

Unit 8 – Syenite to Quartz-Syenite Porphyry: Pink to greyish white when weathered, with 30 to 60 per cent potassium feldspar phenocrysts and megacrysts that are 20 by 2 millimetres on average and occassionally reach 12 by 2 centimetres in size. The matrix is fine grained and consists of feldspar, amphibole and quartz. Within the stock are xenoliths of diorite consisting of 70 per cent feldspar, 15 to 20 per cent amphibole and minor amounts of other minerals.

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WAJOR OXIDE AND MINOR ELEMENT ANALYSIS, CANTIN CREEK; SAMPLES PLOTTED ON FIGURES 1-00-4 AND 1-00-5 SHOWN WITH AN ASTERISK

TABLE 1-20-1

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Note: Values below the analytical detection limit are listed as zero.

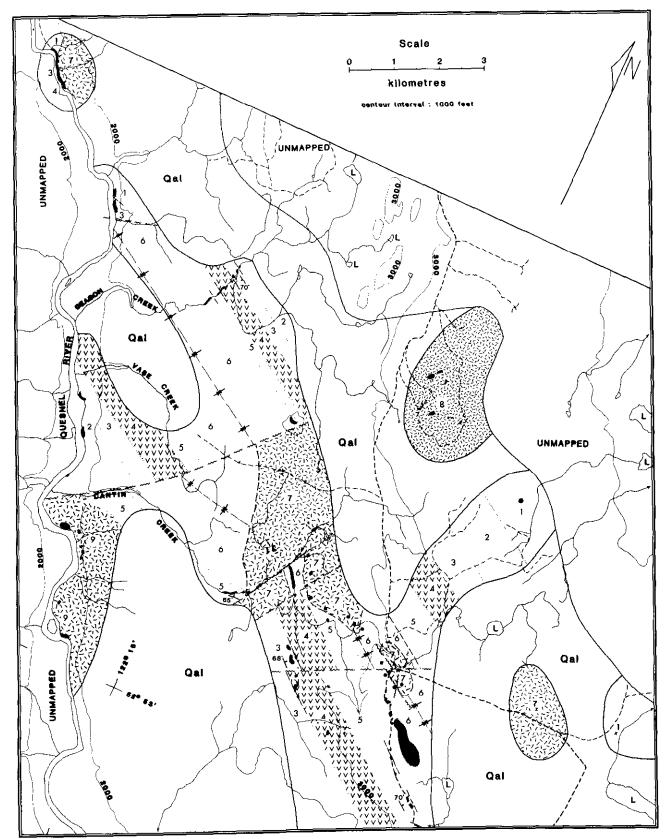


Figure 1-20-3a. Cantin Creek Study Area.

AGE	STRATIGRAPHI	C UNIT #	NA	ME		DESCRIPTION				
- T	COLUMN	Qal	Alluv	rium		lidated till and den, sand to cobble size, sorted				
2		9	Grano	diorite		edium-grained, biotite- , locally carbonate-				
R I		8	Quartz : Syenite F	• •	white, r	pale grey and chalk- medium to coarse-grained, megacrystic				
		7	Gabbro, P Diorite, M Syen	onzonite,	monzonit	tiated diorite to e stock; mafic cumulate ite, in part				
		6	Polylithic feldspat	breccias, hic tuffs	feldspat	with heterolithic clasts; h)c clasts and tuffs contains basaltic detrius				
		5		pathic cke		feldspathic wacke, n by a thin limestone				
, k		4		n Basalt ows		flows, massive, ing pyroxene and feldspar ysts				
		3	-	- bearing cke		crystal - lithic wacke, t pyroxene and feldspar				
		2	Basalt	t flows	grained,	een flows, massive, fine- , containing small r and pyroxene ysts				
Υ Ψ		1	a	llite nd stone	bedded,	≥y to black, thinly locally thin beds of thic wacke				
	, — — ,	Geological contacts defined, approximat	e, assumed	+ +,	∔ · ∔ s	Synform; approximate, assumed				
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<	🤉 , ×	Outcrop; large, small	I	 	L	Limit of mapping				
~	70	Bedding attitude		-d	S	Streams and lakes				
•		Drill hole; diamond or	percussion	2000	C	Contours: elevation in feet				

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Figure 1-20-3b. Legend for Figure 1-20-3a.

TABLE 1-20-2 MINOR ELEMENT ANALYSIS – RANGE AND MEAN VALUE ACCORDING TO MAP UNITS.

Map Unit	No of Samples	Au	Ag	Cu	РЬ	ments in pp Zn	Hg	Sb	As	Ni	Мо
2	l теал	0	0	54	6	133	0	2.5	6	16	3
3	3 range	0-30	0	89-246	4-8	60-93	0-11	3-3.4	0-3	13.21	0-2
	mean	10	0	148	5	80	4	2.1	1	16	1
4	6 range	0-40	0	19-510	8-16	24-97	0-60	0-5.3	1-28	25-221	0-2
	mean	20	0	162	11	77	23	1	12	95	1
5	2 range	0-30	0	8-13	8-14	105	0-10	0.6-1.8	11-14	24-49	0-1
	Ũ	15	0	11	11	105	5	1.2	13	37	1
6	7 range	0-30	0	4-66	5-25	19-153	0-560	0-2.8	2-30	1-165	0-5
	mean	10	0	32	12	65	87	1.2	15	53	2
7	5 range	20-30	0	11-295	0-8	55-98	15-2000	0-0.8	0-30	42-827	0-2
	mean	22	0	125	4	78	419	0.3	13	275	1
8	5 range	0-50	0	3-69	12-28	0-94	0-82	0-6	1-22	8-26	1-1
	mean	24	0	31	17	52	42	2	10	14	4
9	3 range	0	0	2-11	16-27	18-63	0-10	0-1.1	0-14	5-18	0-8
	mean	0	0	6	20	37	3	0.4	9	12	5
All Units	32 mean	15	0	79	11	71		1.2		80	2

Note: Values below the analytical detection limit are shown as zero.

Late-stage syenite dykes, consisting of 80 per cent potassium feldspar, 10 per cent amphibole and 5 per cent quartz cut the stock. Veins of granular white quartz are found throughout the stock and crosscut all phases of the unit.

Unit 9 – Granite Stock: White to greyish white with a porphyritic texture. It is composed of 60 to 70 per cent potassium feldspar, 15 to 20 per cent quartz, 5 to 7 per cent biotite, and 5 to 10 per cent sodium feldspar. Locally the rock has been intensely carbonatized.

STRATIGRAPHY

The age of the map units has been determined by lithologic correlation as no fossil control has been established in the map area.

The age of Unit 1, which is equivalent to Bailey's (1988) Unit 1, has been determined elsewhere (Struik, 1986) to range from Middle Triassic to Late Triassic, mainly on the basis of conodont dating. Units 2 through 5 are stratigraphic equivalents to Bailey's Unit 2 and are thus Late Triassic, probably Norian in age. According to intrusive relationships, Units 6 and 7 are probably Early Jurrasic; Unit 8 appears to be similar to other dated alkalic stocks of Early Jurassic age. However, radiometric data of Panteleyev (1987) and the presence of much hydrothermal alteration suggest a longer period of intrusive activity, possibly well into the Middle Jurassic. Unit 9 is equivalent to Bailey's Unit 9 and so is most probably Cretaceous in age.

STRUCTURAL GEOLOGY

Due to the variability of bedding attitudes and sparse outcrop distribution, it is difficult to interpret details of the regional structure. Based on the established stratigraphic column, the map area is probably underlain by a tight syncline. The southwest limb is relatively well preserved in the southeast and central parts of the map area and trends approximately 130° with dips of 60° to 70° northeast. In the central area, the strata are offset by a fault and are tightly folded and locally overturned. The northeast limb trends approximately

 120° and dips 70° southwest. In the northwest, the intrusion of the granite porphyry (Unit 9) has locally steepened or overturned the strata. The strata are crosscut and offset by northeast to northerly trending normal faults.

PETROCHEMISTRY

Thirty-two rock samples were analyzed for major oxides and minor elements (see Table 1-20-1). X-ray flourescence was used for all major oxides and minor elements, Rb, Sr, Y, Zr, Nb, U, Th, Cr, Ba, Ti, and V. Atomic absorption was used to determine Ag, Cu, Zn, Mo, Ni, As and Sn. Gold was analyzed by fire assay and atomic absorption finish. The data were plotted on a series of discrimination plots to determine the geochemical character of the rocks. To meet the prerequisites of certain diagrams, a number of altered samples with elevated loss on ignition (LOI), H₂O and CO₂ were screened out. Fourteen of the least-altered samples were chosen to be representative of the rock suite. Even this select sample group, when tested by discriminant major-element plots described by Beswick and Soucie (1978) and de Rosen-Spence and Sinclair (1987), reveal that only $A1_2O_3$, SiO_2 , TiO₂, P₂O₅ and possibly Na₂O remain relatively consistent. The other major oxides - K₂O, CaO, Fe₂O₃, MgO and MnO are changed by various degrees. Based on these observations, especially the low TiO₂ content of the rocks, a generalization can be made, as shown on Figure 1-20-4, that the sample suite represents island arc calalkaline basaltic flows deposited in a convergent plate setting. Alkali enrichment noted elsewhere in the Quesnel belt is not as evident in the Cantin Creek rock suite. This is possibly because the more alkalic rocks were not selected for analysis or the analyses were rejected because the alkalic rocks are the most highly altered.

Minor element discriminant plots based on immobile elements are considered to be less affected by alteration. Various plots, some of which are shown on Figure 1-20-5, indicate a volcanic-arc environment of basaltic character. However, the minor element plots are not capable of further resolving whether the magma suites are alkaline or sub-

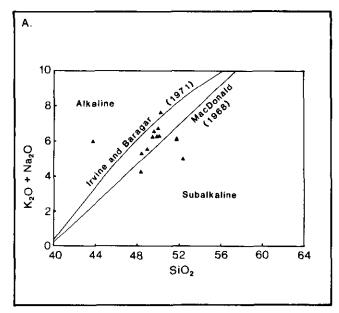


Figure 1-20-4a. Alkalinic-subalkalic divisions according to Irvine and Baragar (1971) and MacDonald (1968).

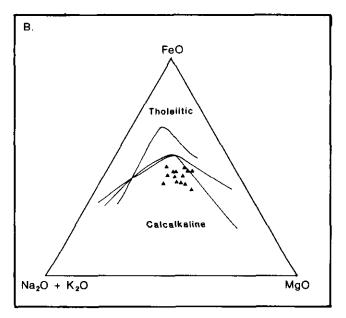


Figure 1-20-4b. AFM diagram after Irvine and Baragar (1971) and MacDonald and Katsura (1964) showing calcalkaline nature of Cantin Creek rocks.

alkaline (tholeiitic) in affinity. These indeterminable minor element data are consistent with other analyses from pyroxene-bearing basalts that were deposited during early Quesnel arc volcanism (Bloodgood, 1987; Morton, 1976; Bailey, 1978).

ALTERATION AND MINERALIZATION

At least five related types of relatively low-temperature hydrothermal alteration or burial metamorphism have been

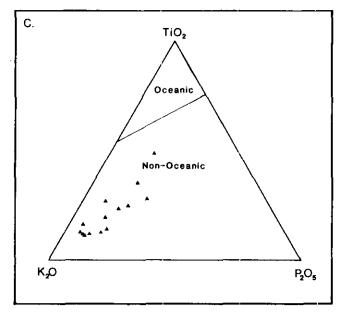


Figure 1-20-4c. TiO_2 -K₂O-P₂O₅ plot after Pearce *et al.* (1975) showing the low TiO_2 and potassic nature of the non-oceanic rocks.

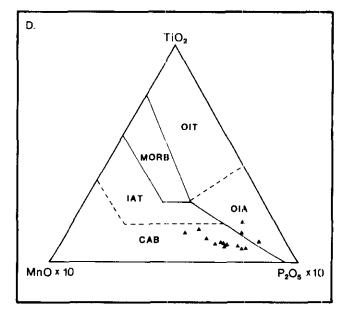


Figure 1-20-4d. Tectonic setting discrimination according to Mullen, (1983) showing calkalkaline (CAB) to ocean island (alkaline) basalt (OIA) character. IAT = island arc (low K) tholeiite, OIT = ocean island tholeiite; MORB = mid-ocean ridge basalt.

identified at Cantin Creek: carbonatization, chloritization, epidotization, pyritization and zeolitization. All types are found as pervasive alteration and in some veins. Carbonatization is most common throughout the map area and is not confined to any specific map unit. Chloritization is also common, but is best developed in Units 5 and 6. Epidote alteration is confined to Unit 6 and pyritization to Unit 5. Zeolite alteration is essentially restricted to the vesicles of flow rocks. Units 5 and 6 are the most pervasively altered, seemingly because of their high porosity.

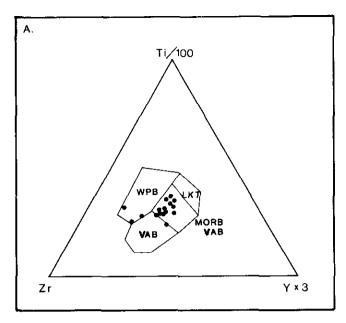


Figure 1-20-5a. Discriminant minor element tectonic setting plot after Pearce and Cann (1973). VAB = volcanic arc basalt; WPB = within plate basalt.

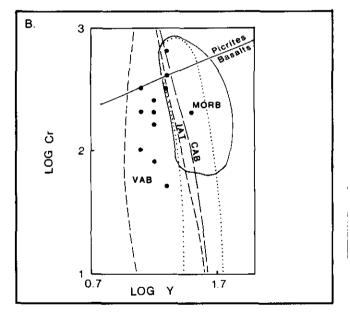


Figure 1-20-5b. Minor element discriminant plots after Pearce (1982).

A series of 21 rock samples were analyzed for gold and related elements (Au, Ag, Cu, Pb, Zn, Hg, Sn, Mo, Ni; *see* Table 1-20-2) in order to study the relationship between alteraton and mineralization. Statistical analysis shows there is no clear association between any specific rock unit and gold enrichment. Furthermore, there is no apparent correlation between alteration type and gold. However, comparing Cantin Creek to the Horsefly area (Morton, 1976) and the Nicola Group (Preto, 1979), the data are significantly different. In the Horsefly area, the maximum nickel, copper and zinc values are 128; 198; and 125 ppm; in the Cantin Creek area these values are much higher at 877; 510; and 187 ppm respectively. In Nicola rocks, the mean values of nickel for

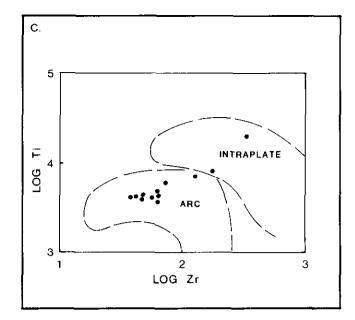


Figure 1-20-5c. Minor element discriminant plots after Pearce (1982).

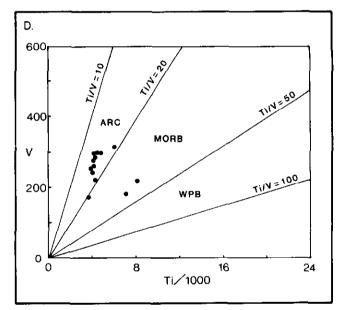


Figure 1-20-5d. Ti-V discriminant plot after Shervais (1982) distinguishing tectonic settings.

flow rocks, tuffs and intrusive rocks are approximately 17; 10; and 6 ppm; for copper, 77; 56; and 45 ppm; for zinc, 93; 89; and 76 ppm; and for lead, 12; 10; and 6 ppm. In contrast, the mean values for the same rocks in the Cantin Creek area are: nickel, 95; 53; and 114 ppm; copper, 162; 32; and 61 ppm; zinc, 77; 65; and 59 ppm; and lead, 11; 12; and 13 ppm. These values clearly indicate that the copper and nickel content in Cantin Creek rocks is greater than that of Nicola Group rocks in general. These data demonstrate the basic nature of the basal Quesnel volcanic units and their pyroxenerich erosional products compared to the more differentiated Nicola successions. Gold, silver and related-element data from the map area cannot be compared with other areas in the

Quesnel belt due to the limited data available. However, with mercury values up to 2 ppm, mean gold values of 15 ppb and up to 50 ppb, the potential for significant gold concentrations in the area is indicated.

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