

# British Columbia Geological Survey Geological Fieldwork 1995

## NATURAL ACIDIC DRAINAGE IN NORTHERN VANCOUVER ISLAND -ITS PLACE IN GEOENVIRONMENTAL ORE DEPOSIT MODELS

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*KEYWORDS:* acid rock drainage, pH, dissolved metals, northern Vancouver Island, geoenvironmental ore deposit models, acidic water geochemistry, advanced argillic alteration.

#### INTRODUCTION

During the course of mineral deposits studies in Northern Vancouver Island (Figure 1) as part of an integrated project (Panteleyev *et al.*, 1994), stream waters emanating from, and flowing through, large areas of pyritic rocks were found to be strongly acidic (commonly <4 pH). The bedrock sources for the acidic waters are hydrothermally altered and mineralized Bonanza volcanic rocks and some Island intrusions containing porphyry copper mineralization and large zones of advanced argillic alteration.

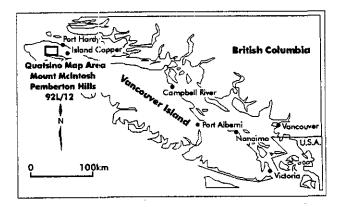


Figure 1. Location of water sampling in Quatsino map area (NTS 92L/12).

### GEOLOGICAL SETTING AND MINERALIZATION

The regional geological setting of the Quatsino Sound area is described by Nixon *et al.* (1994) and Hammack *et al.* (1994). The style of mineralization and alteration 15 to 40 kilometres to the west of Island Copper mine, mainly in the Bonanza volcanic rocks and to a lesser extent in the Island intrusions, is discussed by Panteleyev and Koyanagi (1993, 1994). Ages of hostrocks, hydrothermal systems and alteration minerals are summarized by Archibald and Nixon (1995) and Panteleyev *et al.* (1995).

The principal mineral deposits in the study area are the intrusion-related, pyrite-rich Hushamu and Red Dog porphyry copper deposits. A number of related, smaller pyritic zones are also present, for example the Hep deposit. The Hushamu deposit has an overprint of advanced argillic alteration that is expressed as the siliceous "capping" on Mount McIntosh. The alteration (quartz-pyrophyllite-kaolinite-alunite-zunyite-diaspore) is superimposed, together with vuggy quartz, and weakly developed, high-sulphidation epithermal mineralization, on the deeper, underlying Hushamu purphyry copper deposit. The advanced argillic overprinting occurred when the late hydrothermal system collapsed onto the predominantly copper-bearing, quartzdeeper sericite/illite-pyrite and quartz-amphibole-magnetitechlorite alteration zones. Both the upper Mount McIntosh and the underlying Hushamu zones are cut by younger, siliceous hydrothermal breccias. The weakly developed epithermal mineralization that accompanies the advanced argillic and vuggy quartz alteration contains abundant pyrite with minor enargite, chalcopyrite, covellite and chalcocite (Pantelevev and Koyanagi, 1994; Dasler et al., 1995).

There is widespread advanced argillic alteration to the east of the Hushamu deposit throughout the Pemberton Hills over a distance of about 11 kilometres, and in the western part of the study area at the Rec Dog deposit and the nearby Northwest Expo claim area It is evident as pyritic quartz-kaolinite-alunite alteration in the rhyolitic Bonanaza map units. Centres of most intense alteration seem to be associated with rhyolite dome emplacement, associated vent deposits and hydrothermal breccia bodies. In addition to abundant pyrite, quartz-alunite is widespread and there is rare enargite, covellite and chalcocite; native sulphur is locally abundant. The strongly altered, siliceous, aluminous rocks have been hydrothermally acid leached. They lack buffering (acid neutralizing) capacity and weathering of the highly pyritic rocks gives rise to extensive zones of near-surface, secondary (supergene) acid-leached limonitic rocks.

#### **GEOCHEMICAL WATER SAMPLING**

Starting in 1991, and continuing in 1992 and 1993, stream and various standing waters were measured directly in the field for pH, conductivity and total dissolved solids (TDS) at 248 sites, using a Corning Checkmate<sup>TM</sup> M90 portable microprocessor. Preliminary results are summarized by Koyanagi and Panteleyev (1993, 1994). Water geochemical samples were taken

Sample	E	Cond	TD6	UTM Zone 10	one 10	Location	<b>S04</b>	Fe	Si	V	Mg	Ca	Mn	Cu	Ρb	Zn	Ba	°C	ž	٩
No. 1991 semples		ŧ.	lm/B	Easting	Northing				5											
EC91AP-16	3.9	163	82	683092	5609526	Hushamu Creek	31	0.01	6.8	1.9	1.7	5.2	0.11	0.01	0.02	0.01			< 0.01	0.05
EC91AP-19	3.8	177	92	683209	6611889	Hushamu Ck. tributary	39	0.01	8.7	1.9	2.2	4.0	0.12	0.00	0.01	0.01			<0.01	0.03
EC91AP-20	4.3	120	61	582996	5612577	Hushamu Creek	18	0.01	6.9	1.0	1.9	4.7	0.07	0.01	0.0	0.01			<0.01	0.01
EC91AP-22	3.5	244	126	581785	5613442	Hushamu Ck. tributary	50	0.21	6.9	6.3	1,4	0.2	0.30	0.01	0.02	0.01			<0.01	0.05
	4,3	141	71	581451	5613942	Hushamu Lake	32	0.01	5.8	1.6		6.4	0.10	0.02	0.02	0.01		v	< 0.01	0.03
EC91AP-24	3.7	64	27	581488	6614083	Hushamu Ck. tributary	228	0.55	21.5	16.5		23.3	0.47	0.32	0.05	0.04			0.04	0.09
	6.2	51	25	582901	5609635	Hushamu Ck. tributary		0.10	1.9	0.1	0.6	2.1	0.01	0.00	0.01	0.03			<0.01	0.02
EC91AP-30	4.1	<u>10</u>	50	679683	5614349	Hepter Ck. tributary	18	0.01	3.7	0.8	0.9	3.0	0.30	0.07	0.01	0.01			<0.01	0.04
	3.9	119	62	679689	6614265	Hepler Ck. tributery	20	0.07	3.7	1.2	1.2	3.2	0.06	0.00	0.01	0.01			<0.01	0.04
EC91AP-34	5.6	40	20	578313	5612787	Goodspeed R. tributary	e	0.01	1.4	0.2	0.5	1.5	0.01	0.00	0.00	0.03			0.01	0.01
EC91AP-35	4.8	15	œ	680888	5613535	Mount McIntosh	2	0.29	0.3	0.3	0.1	0.0	0.01	0.00	0.01	0.01			< 0.01	0.01
	3.6	88	45	580956	5612541	Mount Meintosh	14	0.56	0.7	1.0	0.8	0.0	0.01	0.0	0.00	0.01			<0.01	0.01
	3.9	159	79	581612	5610201	South McIntosh	29	0.01	5.4	1.9	1.7	6.9	0.25	0.00	0.01	0.01	0.03 <	<0.01 <	0.01	0.01
1992 samples																				
EC92AP-1	5.6	76	38	577726	5611934	Goodspeed R. tributary	13	0.13	3.9	0.2	1.2	0.0	< 0.01	0.01	0.01	0.03				<0.01
EC92AP-3	5.6	23	12	583032	5615140	Mead Creek	5	0.41	2.2	0.5	0.2		< 0.01	0.00	0.01	0.04				<0.01
EC92AP-6	<b>6</b> .5	35	18	584095	5616670	Mead Creek	9	0,19	2.9	0.2	0.5		<0.01	0.01	0.01	0.11				<0.01
EC92AP-10	5,0	74	36	579235	5617182	Hepler Creek	26	0.45	5.3	0.5	1.2	5.7	0.04	0.01	0.01	0.01				<0.01
EC92AP-13	3,9	152	75	578378	5615969	Hepler Ck. tributary	51	0.35	7.8	3.5	1:1		0.02	0.03	0.01	0.08				0.02
EC92AP-15	4.7	51	26	583015	5612783	Huehamu Ck. tributary	17	0.29	4.5	0.3	0.6		0.02	0.00	0.01	0.02				< 0.01
EC92AP-16	4.4	151	75	582246	5613217	Mount McIntosh	54	9.20	8.5	1.6	2.8	7.6	0.07	0.02	0.01	0.02				<0.01
EC92AP-18	3.2	269	128	580958	6612530	Mount McIntosh	58	0.14	3.3	0.1	0.7		0.02	0.01	0.01	0.06				0.02
EC92AP-19	5.4	55	27	579806	6611551	Cleeklagh Creek	₽	2.30	1.1	2.9			<0.01	0.01	0.01	0.01				<0.01
_			63	680697	5609698	Cleeklagh Creek	72	0.68	8.2	3.2			0.39	0.01	0,01	0.03	v		< 0.01	0.02
		~	1190	580439	5611893	South McIntosh ditch	1300	88.80	36,6	45.7			10.00	0.15	0.11	0.65				0.3
	3.6	219	105	582100	5610975	South McIntosh	67	0.29	5.4	1.7			0.79	0.01	0.01	0.02				<0.01
	4.8	53	25	584759	5608931	Hushamu Ck. tributary	12	2.98	3.5	9.0			0.08	8.0	0.01	0.01				<pre>&lt; 0.01</pre>
	ی ۲	252	124	585996	5609231	Youghpan Creek	72	2.23	6.9	3.5	н. Н		0.18	10.0	5.0	20.0			10.02	55
	2.9	311	157	585514	5609187	Youghpan Ck. tributary	113	5.60	11.3	0.4	2.1	14.7	0.44	0.01	0.0	0.03				20.02
	6.3 •	8	28	585853	6611835 500000	N. Youghpan Creek	1	0.17	2.9	5 F	 	3.0 1	0.06	0.0	5.0	0.01			5.00	10.02
	4 2 0	166	63	582226	16/9099	Mouth Youghpan Creek	60	69.0	а. 4 с	> r - <	ה - כ	2.7	07.0	50	10.0	70.0				36
	0.4 0.4	16 17	43	586391 505654	01680095	Youghpan CK. tributary Vousitions Ct. 4-thutani	2 9	17.1 9 2E	9 V 0 P		8.0 F	o u v v	0.0	50	0.0	0.00				000
ECSZAT-00	+ U 7 C	26AU	8 I I 8	000304	5608190	Youghpan Ck. Moutary Vouchaan Ck. telbutary	88		ς α - Γ	t a N a		9.0	0.16	000	600	0.03			<0.01	0.04
	2	146	2 2	588699	5607854	Youghpan Ck. tributary	49	1.12	6.7	0.4	- <del>-</del>	16.4	0.13	0.01	0.01	0.01				< 0.01
		•				• • • •														
	3.6	206	102	572115	5618055	Red Dog	62	0.24	9.0	3.2	3.4	5,4	0.28	0.38	0.01	0.02				0.02
S493VK-02	6.4	76	38	571641	5618113	Red Dog	23	0.39	6.6	0.1	6.1	7.8	0.01	0.01	0.0	0.00				<0.01
S493VK-03	5.2	88	43	671690	6617375	Red Dag	23	0.49	13.2	5.4	3,8	28.7	0.49	0.02	0.02	0.04				0.03
S493VK-04	3.2	471	240	580731	5611275	Clesklagh Ck. tributary	150	1.30	6.2	0.5		7.2	0.04	0.09	0.01	0.0				<0.01
S493VK-05	4.1	230	120	580630	5609747	Cleskiagh Creek	20	0.52	8.6	2.9	2.4	14.6	0.36	0.01	0.02	0.02				0.02
S493VK-06	3.4	158	80	583092	5609526	Hushamu Creek	61	0.46	8.7	2.0		7.9	0.12	0.02	0.03	0.02				< 0.01
	3.8	436	219	582069	5610698	South McIntosh	149	0.21	14.2	6.7			0.72	0.01	0.03	0.03				0.06
	6,5	73	35	579603	5614379	Hepter Ck. swamp	76	1.58	5.1	0.4			0.02	0.02	0.02	0.00				<0.01
	8	190	96	679235	5617182	Hepler Creek, S. fork	58	1,44	8.0	1.8			0.14	0.01	0.03	0.01				<0.01
_		1017	627	58007	5614325	Hushamu deposit ditch	338	2.40	14.8	20.1	6 6		0.43	1.37	0,08	0.16				0.12
	4.3	137	20	582034	5613403	Hushamu Ck. headwater	51	1.35	6.9	1.3	2.9		0.09	0.03	0.03	0.02				
	5.2	58	29	582993	5612749	Hushamu Creek	14	0.38	4.4	0.3	0. F	3.7	0.09	0.01	0.01	0.01				0.01 20.01
S493VK-13	3.7	181	06	683227	5611897	Hushamu Ck. tributary	58	1.48	11.0	2.1	2.9	0.7	0,16	0.01	0.03	10.0	v 07	> 10'0 >	× 10.0 ×	10.0

TABLE 1: WATER CHEMISTRY - MOUNT MGINTOSH/PEMBERTON HILLS AREA, NORTHERN VANCOUVER ISLAND, IN ppm

Sample	ΡH	Cond	TDS	UTM Zone 10	one 10	Location	<b>S04</b>	Fe	ŝ	₹	Mg	٦	ЧW	C.	£	Zn	Ba	రి	ž	ð
No.		8.7	[m/8	Easting	Northing															
1993 samples (continued)	s (cont	(penuj	1				2	ļ	(		4		0000		2.05	50.0		÷0 0	500	100
S493VK-14	3.9	215	88	581533	6610239	Clesklagh Ck. tributary	61	0.47	6.8	8.	5.S	10.4	0.28	20.0	60.0 0	10.0	61.0	5		
S493VK-15	ດ ເດິ	65	32	555572	5611986	top   600 Road	r. *-	4.03	3.0	e) ()	••	e) ej	0.57	0.00	0.01	80	0.04	5.5		
S493VK-16	5.4	66	28	585896	5611834	top H600 Road	12	0.15	3.7	0.2	1.0	2.7	0.08	0.01	0.02	0.00	0.05	<0.01	<0.01	v
S493VK-17	3.5	408	203	585165	5609947	Youghpan Creek, head	138	1.21	15.0	5,9	4.5	20.1	0.51	0.02	0.06	0.05	0.03	0.02		
S493VK-18	3.3	449	232	585523	5609163	Youghpan Creek	126	2.53	10.6	3.9	3.4	13.9	0.42	0.02	0.06	0.04	<0.01	0.02	-	
S493VK-19	3.7	250	130	586195	5609062	Youghpan Creek	65	2.04	6.4	3.2	1.8	10.8	0.17	0.02	0.04	0.02	< 0.01	0.01	<0.01	
S493VK-20	4.2	176	88	582226	5606791	Youghpan Ck. at Main	65	0.41	5.6	1.5	1.8	13.0	0.19	0.01	0.03	0.01	<0.01	< 0.01	<0.01	
S493VK-21	6.4	189	96	593651	5606900	Wakalish Ck. mouth	18	0.01	2.9	0.1	2.8	4.7	0.01	0.01	0.02	0.00	0.02	< 0.01	<0.01	<0.01
1995 samples																				
EC95-01	6.2			591015	5606839	Wanokana R. at Main	:	0.39	3.3	0.1	0,7	3.5	0.01	0.01	0.0	0.00	<0.01	<0.01	<0.01	
EC95-02	4 5			582226	5606791	Youchoan Ck. at Main	56	0.52	6.3	1.1	1.5	13.6	0.15	0.01	0.02	0.01	0.05	0.01	< 0.01	< 0.01
ECO5.03				584949	5608906	H600 Road at Main	σ	0.99	3.8	<.0.1	1.0	6.0	0.04	0.01	0.03	0.10	0.01	< 0.01	< 0.01	< 0.01
ECOE-OA				583093	ERNOF 26	Hushamy Ck at Main	43	0.75	7.4	0	1.7		0.09	0.02	0.06	0.02	0.03	0.01	< 0.01	< 0.01
	4 P			582069	EGIOROB	H1000 Boad S McIntosh	200	4 00	141	4.6	4.2		0.43	0.01	0.31	0.03	0.04	0.01	0.01	
				501437	EGIORG1	Clasticath Ck tributero	5	0 54			5.0		0.25	0.01	00.0	0.01	0.04	< 0.01	< 0.01	
	n o f c			50142V	561127E	Contract CK: structure	5	414	о 1		i e		0.39	000	0.21	0.03	0.04	0.0	<0.01	
EC95-07	0 0 7 9			101000	00112100	Managii CK at CLISO	2						510	500	010	100	1002	002	<0.05	
EC35-08	N 1			6678/G	201/100		5 9	5			) ( ) (								1007	
EC95-09	4			561451	5613942	Hushamu Lake	7 C C	4 8 5 6	0 + 0 r	2	) • •				5 6					
EC96-10	÷			<b>582996</b>	P61 Zb //	Hushamu CK, headwater	Ŗ	3	-				5	20.0	3		5			
1995 samples were filtered and acidified; 1992 and 1993 Analytical method: ICP. Laboratory: 1991 - MinEn; 1992,	s were thod: 1	filtered CP. La	l and ac borator	sidified; 199 y: 1991 - N	32 and 1993 AinEn; 1992,	i samples were acidified but not fittered; 1991 samples were untreated. , 1993 - Eco-Tech; 1995 - CanTech.  iCP data for other elements are available, upon request, from the senior author.	t not filt CanTecl	i iCP o	991 sai lata for	other e	/ere uni ilement:	reated.	/ailablo,	npan re	iquest, 1	from the	senior al	uthor.		
TABLE 3:	REF	<b>LIC</b>	VTE V	VATER	SAMPLE	TABLE 3: REPLICATE WATER SAMPLES, NORTHERN VANCOUVER ISLAND, IN ppm	NCOL	IVER	ISLA	ND, I	N ppr	E								
Sample	Ha	Cond	TDS	UTM 2	UTM Zone 10	Location	<b>S04</b>	Fe	Si	N	Mg	Ca	Mh	Сu	£	ΨZ	Ba	ပိ	Ż	ž
No.		87	1m/0	۳.	Northing															
EC95-02	4.5			582226	5606791	Youghpan Ck. at Main	56	0.52	6.3	1.1	1.5	13.6	0.15	0.01	0.10	0.01	0.05	0.01	<0.01	v
EC92AP-33	4.2	166	83	582226	5606791	Mouth Youghpan Creek	69	0.89	8.4	1.7	1.9		0.26	0.01	0.01	0.02	0.09	< 0.01	<0.01	
S493VK-20	4.2	176	88	582226	5606791	Youghpan Ck. at Main	65	0.41	5.6	1.5	1.8	13.0	0.19	0.01	0.03	0.01	<0.01	<0.01	<0.01	<0.01
FC96-04	4.2			583092	5609526	Hushamu Ck. at Main	43	0.76	7.4	1.0	1.7	7.1	0.09	0.02	0.14	0.02	0.03	0.01	<0.01	< 0.01
EC91AP-16	9.6	163	82	583092	5609526	Hushamu Creek	31	0.01	6.8	1.9	1.7	5.2	0.11	0.01	0.02	0.01	0.03		< 0.01	
S493VK-06	3.4	158	8	583092	5609526	Hushamu Creek	61	0.46	8.7	2.0	2.2	7.9	0.12	0.02	0.03	0.02	<0.01	<0.01	<0.01	< 0.01
FC95-05	3.7			582069	5610698	H1000 Road, S. McIntosh	100	4.00	14.1	4.6	4.2	17.6	0.43	0.01	0.39	0.03	0.04	0.01	¢0.01	< 0.01
S493VK-07	3.8	436	219	582069	5610698	South McIntosh	149	0.21	14.2	5.7	4.2	32.2	0.72	0.01	0.03	0.03	0.09	0.01	< 0.01	0.06
EC95-08	4.2			579235	5617182	Hepier Creek	54	1.61	7.4	1.1	2.0	9,9	0.11	0.01	0.18	0.01	< 0.01	< 0.01	< 0.01	
S493VK-09	3.8	190	96	579235	5617182	Hepier Creek, S. fork	58	1.44	8.0	1.8	2.5	11.0	0,14	0.01	0.03	0.01	0.04	<0.01	< 0.01	<0.01
FC95-09 FC91AP-23	4.7	141	5	581451 581451	5613942 5613942	Hushamu Lake Hushamu Lake	<b>49</b> 32	0.44 0.01	9.8 9.8	1.7 1.6	3.0 2.3	10.4 6.4	0.11 0.10	0.05 0.02	0.08 0.02	0.04 0.01	0.01 0.01	0.01 0.01	<0.01 <0.01	<0.01 0.03
				I																

1996 samples were filtered and acidified; 1993 and 1992 samples were acidified but nut filtered; 1991 samples were untreated. Analytical method: ICF. Laboratory. 1991 - MillEll, 1992, 1993, 1994 - Eulw-Tech. ICP date for other elements an available, upon requect, from the confer author.

<0.01 0.01

<0.01 <

<0.01 <0.01

0.04 0.01

0.02 0.01

0.06 0.00

0.02 0.01

0.06 0.07

71

2.1 1.9

0.7

7.1 5.9

0.80 0.01

38 18

5612577 Hushamu Ck. headwater 5612577 Hushamu Creek

582996 582996

61

120

EC95-10 EC91AP-20

4.3 4.6 4.3

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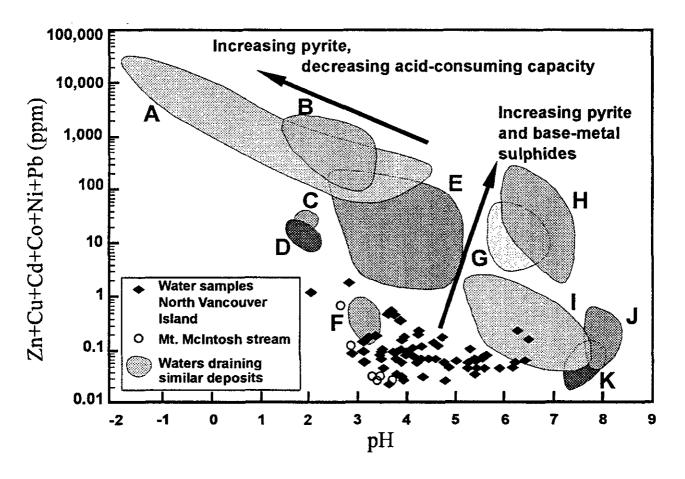


Figure 2. 'Ficklin' diagram of pH and dissolved base metal contents of waters draining diverse mine and mineral deposit types, from Plumlee *et al.* (1994). Fields are: A - pyrite-rich massive sulphides; B - sulphide-rich ores in silica-alunite-kaolinite-clay altered rocks (advanced argillic alteration); C - high-sulphide, low- base metal hotspring ores in acid-altered wallrocks; D - high-sulphide, low-base metal porphyry Mo ores in igneous wallrocks; E - pyrite and base metal-rich polymetallic veins and disseminations in wallrocks with low acid-buffering capability; F - pyrite-rich, base metal-poor veins and disseminations in wallrocks with low acidbuffering capacity; G - pyrite and base metal-rich polymetallic veins with carbonate gangue or carbonate-altered wallrocks; H - pyrite and base metal-rich, polymetallic replacements and veins in carbonate-rich sediments; I - polymetallic veins with moderate to low pyrite and base metal content with carbonate gangue and/or carbonate wallrock alteration; J - pyrite-poor polymetallic replacements in carbonate-rich sediments; K - pyrite-poor, Au-Te veins and breccias with carbonate gangue.

from 55 of the sites and analyzed for their SO<sub>4</sub> and metal contents. Two water samples were submitted; one for SO<sub>4</sub> and the other, acidified with nitric acid, for dissolved metals. In 1995, ten additional water samples were taken for analysis and the pH measured *in situ* using a Cole Parmer model 5941-00 pocket digital pH meter. All the geochemical analytical results are summarized in Tables 1 and 2.

Water samples taken in 1991 were neither acidified nor filtered. Replicate sampling of some sites in subsequent years with acidified samples shows no appreciable differences between the acidified and nonacidified samples for most metals. However the unacidified waters do appear to have lower iron contents and inconsistent lead values compared to acidified samples taken in later years. The 1995 sample suite provided both acidified and filtered samples and used acid-washed, ultra-clean sample bottles. For most metals there is no evident difference between metal content of these most recent, filtered samples and the earlier unfiltered ones. All waters analyzed contain very little dissolved ore metals though iron, and other metals derived from silicate mineral breakdown such as aluminum, calcium and magnesium, are present in abundance. Results of replicated sampling are given in Table 3.

## GEOENVIRONMENTAL ORE DEPOSIT MODEL

Geochemical classification and description of metal distributions within, and in waters derived from, various mineralized environments and ore deposits is currently being undertaken and is referred to as 'geoenvironmental' mineral deposit modeling. The work of the United States Geological Survey, by Ficklin et al. (1992) and Plumlee et al. (1992, 1994), among others, has documented and classified a large number of mine and natural drainages derived and evolved from various mineralized settings. The information is presented graphically in what has been referred to as a 'Ficklin' plot (Figure 2).

The data from the measured and analyzed natural drainages in this study on Vancouver Island are illustrated on Figure 2. Numerous samples have low pH (pH 2 to 4), elevated iron, magnesium, manganese and aluminum, but low levels of copper and zinc (tens to hundreds of ppb) compared to mine drainages. The waters all display low levels of metal loading, rarely greater than 3 ppm combined Zn+Cu+Cd+Co+Ni+Pb.

The effects of dilution are described by Sibbick and Laurus (1995). Their detailed follow-up work in a stream undisturbed by logging in the the Mount McIntosh area, shows that stream acidity and dissolved metals have their sources, and are most concentrated, in the strongly altered rocks from which the streams emanate or pass over. Downstream from the source areas, metal concentrations and acidity decrease due to dilution by tributary streams. Their results are shown on Figure 2 as open circles. The most acidic samples are at the source of the stream in the altered and mineralized rocks; the other circles are sample sites at 100 to 200-metre intervals downstream, where progressively greater dilution effects are evident over the measured distance of 800 metres.

## CONCLUSIONS

The main factors determining the pH and metal contents of drainages are, according to Plumlee et al. (1992): the acid-buffering capacity of the gangue minerals and hostrocks in which the mineralization occurs; the types and abundances of metal-bearing sulphide minerals present; the availability of sulphides for weathering; and the availability of dissolved oxygen during sulphide weathering and drainage generation. Geologic factors influencing the stream geochemistry are: deposit size, hostrock composition, wallrock alteration, surrounding geological and surficial terrains, nature of ore, trace element geochemistry, ore and gangue mineralogy and zonation, mineral characteristics, secondary mineralogy, topography and physiography, hydrology and degree of surface disruption or exploitation.

The natural acidity in the northern Vancouver Island study area is due to the abundance of pyrite in the extensive zones of advanced argillic alteration and the locally underlying porphyry copper mineralization. There is a lack of neutralizing capability in the hydrothermally altered, rhyolitic Bonanza volcanic rocks, and their related granitic bodies. Pyrite, mainly as fine-grained disseminations and fracture fillings, is abundant but little ore metal is present in the acidic leaching solutions, due to an apparent lack of ore minerals in the near-surface (weathering) environment.

In the generally wet coastal climate, rainwater with pH of about 5.6, and the augmentation of streams by surface, groundwater and tributary stream additions, results in downstream dilution of pH and metal contents. Nevertheless, streams maintain their acid character (pH

<5) over distances of a few kilometres from the sources of acidity unless they flow across neutralizing rocks. In the study area, the most effective neutralization takes place in sedimentary rocks of the Parson Bay and Quatsino limestone map units and the weakly altered, zeolite and calcite-bearing basaltic units of the Bonanza volcanics.

The sources of ferruginous, base metal deficient acidic drainages derived from hydrothermally ac.dleached rocks of northern Vancouver Island can be classified as 'pyrite-rich, base metal-poor veins and disseminations in wallrocks with low acid-buffering capacity' (category F on Figure 2), or more simply, (an) "acid, low-metal environment".

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