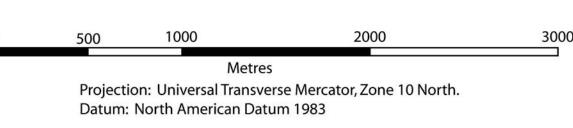




QUATERNARY GEOLOGY OF RICHMOND, **BRITISH COLUMBIA**

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Scale 1:20 000 (when printed on 36x60 inch paper)

This map shows geological variation in the Quaternary deposits underlying the City of Richmond, British Columbia, and has been prepared as part of an earthquake hazard mapping project in the city. Richmond is located in one of the most seismically active regions in Canada (Rogers, 1998), and the deltaic sediments on which the city is built can amplify ground motions and are susceptible to earthquake induced liquefaction (e.g. Byrne, 1978; Finn et al., 1989; Byrne and Anderson, 1991; Clague et al., 1992, 1997, 1998b; Watts et al., 1992; Rogers et al., 1998; Harris et al., 1995, 1998; Cassidy et al., 1997; Cassidy and Rogers, 1999; Levson and Monahan, 2004). Because the variation in earthquake hazards is most affected by geological conditions in the upper 20 to 30 metres, this map primarily reflects geological variation in the upper 20 to 30 metres of the Quaternary

Quaternary Geology Framework

Richmond is located entirely within the Fraser River delta, which is the largest delta in western Canada. The deltaic deposits are entirely Holocene in age and overlie Late Pleistocene glaciogenic sediments (Clague et al., 1983, 1991, 1998a; Luternauer et al., 1994). The thickness of the deltaic deposits underlying Richmond varies from less than 20 metres locally along the North Arm of the Fraser River to a maximum known thickness of 305 m at Richmond City Hall (Dallimore et al., 1996). Deltaic deposits are generally thicker than 50 m at distances greater than 1 km from the delta margin, and the northern margin is located immediately north of the North Arm in the southernmost parts of Vancouver and Burnaby. However, a buried ridge of Pleistocene sediments, above which deltaic deposits are locally less than 20 metres thick, trends southeast across the city from the western edge of Mitchell Island, which is located in the North Arm of the River

Monahan et al. (1997, 2000a), Claque et al. (1998a), and Monahan (1999) provide descriptions of the stratigraphy of the delta. The deltaic section can be subdivided into topset, foreset, and bottomset units. Topset deposits were deposited on the delta plain in river channel, tidal flat and floodplain environments, and consist of sand, silt and locally peat. The topset forms the uppermost 20 to 35 metres of the deltaic sequence, so that this map primarily reflects geological variations in the topset, which are discussed in more detail below. Foreset deposits underlie the topset sequence and were deposited on the delta slope beyond the river mouth. They are up to 165 metres thick, consist of sands and silt interbedded on a variety of scales, and are characterized by seaward dips that are commonly 70 near the top of the foreset. The modern equivalents of the foreset are currently being deposited on the delta slope offshore of Sand Heads. Bottomset deposits were deposited beyond the delta slope and consist of mainly clayey silt. They occur only in deeper deltaic sections, where they are up to 120 metres thick. Pleistocene deposits underlying the delta generally consist of tills and other sediments that have been overridden by glaciers and are overconsolidated. However, in the deeper deltaic sections, a few metres of normally consolidated Late Pleistocene glaciomarine clay and silt occur between the deltaic sediments and overconsolidated Pleistocene deposits.

Fraser Delta Topset Deposits

In Richmond, the topset deposits thin from 35 metres in eastern Lulu Island, near the apex of the delta, to 20 metres or less at the sea dykes along western Lulu Island (Claque et al., 1983, 1998a; Williams and Roberts, 1989; Monahan et al., 1993, 1997; Monahan, 1999). The westerly thinning of the topset is the result of a relative sea level rise of 13 metres during the Holocene, most of which occurred between 8000 and 4500 years B.P. (Clague et al., 1983; Williams and Roberts, 1989, 1990). The topset forms an overall fining-upward sequence that grades up from a lower sand facies, through an interbedded sand and silt facies, to an upper organic silt facies. The latter is locally overlain by peat.

The topset is dominated by the lower sand facies, which is generally 8 to 30 metres thick. It has an erosional base with several metres of local relief, and is commonly organized into one or more decametre scale fining upward sequences. It has been interpreted to represent a complex of distributary channel deposits (Monahan et al., 1993, 1995, 1997; Hutchinson et al., 1995; Monahan, 1999).

Two subfacies of the sand facies occur in Richmond and with the overlying interbedded sand and silt facies form two distinct facies successions in the topset (Monahan et al, 1993, 1995, 1997; Clague et al., 1998a, Monahan, 1999). In the first facies succession, the sand facies includes shell debris, particularly near the base. The interbedded sand and silt facies is up to 5 m thick, extensively bioturbated, also shell-bearing and contains an intertidal foraminiferal fauna (Williams, 1988). This facies succession is interpreted to represent sands deposited in migrating distributary channels in a tidal flat environment and capped by tidal flat deposits (see Clague et al., 1983). The shell debris was incorporated into the distributary channel sands as channels migrated and eroded earlier tidal flat deposits. In the second facies succession, the sand facies does not contain shell debris and includes sands that are slightly coarser than in the first succession. The overlying interbedded sand and silt facies is generally only a few metres thick, well-bedded and consists of thinly interbedded fine to medium sand and laminated silt. The latter facies succession is interpreted to represent sands deposited in distributary channels in a delta floodplain environment, capped by silts and sands deposited in active bar tops and slough environments during channel abandonment.

Locally, the interbedded sand and silt facies thickens at the expense of the underlying sand facies to form narrow curvilinear bodies that are up to 20 metres thick. These deposits (represented by map unit szF) are commonly associated with historic sloughs on the delta plain (Township of Richmond, undated) and are interpreted to represent the fill of partially abandoned channels. They range in scale from those that are a few tens of metres across and can be traced 100 to 200 metres, to one along the Main Channel that is 700 metres wide and can be traced approximately 4 kilometres.

The interbedded sand and silt facies is overlain by an organic silt facies consisting of laminated and organic clayey silts deposited in the tidal marsh and floodplain environments. Due to the mid Holocene rise in sea level, the organic silts are up to 12 metres thick in eastern Lulu Island, near the head of the delta, and thin westward to less than 1 metre at the sea dykes on western Lulu Island. The organic silts are in turn overlain by peat in parts of central and eastern Lulu Island (Clague et al., 1983; Williams and Roberts, 1989).

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