Are Cordilleran carbonatite hosts of Ta, Nb and REE from the deep mantle?

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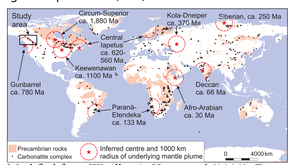
Recommended citation:
Rukhlov, A.S., Aspler, L.B., and Gabites, J., 2019. Are Cordilleran carbonatite hosts of Ta, Nb and REE from the deep mantle? British Columbia Ministry of Energy, Mines and Petroleum Resources, British Columbia Geological Survey GeoFile 2019-08.

BRITISH COLUMBIA | Ministry of Energy, Mines and Petroleum Resources

Introduction

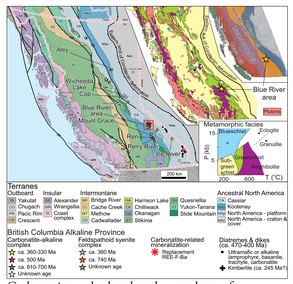
Carbonatites are important exploration targets for Ta, Nb, REE and other commodities. In the Canadian Cordillera, carbonatites were emplaced at ca. 810-700, 500, and 360-330 Ma, forming part of the British Columbia alkaline province. The most prolific Late Paleozoic carbonatites, hosting Nb-Ta and REE deposits, are unusual. In contrast to most carbonatites globally, which are found in intracratonic regions in association with flood basalts, they formed in a more active setting, along the western margin of Laurentia. We investigate metallogeny of these rocks.

Carbonatites and examples of related large igneous provinces (LIPs)



Modified from Woolley & Kjarsgaard (2008), Ernst & Bell (2010), and Ernst & Bleeker (2010).

Carbonatite-alkaline complexes and dike-diatreme swarms defining the British Columbia Alkaline Province

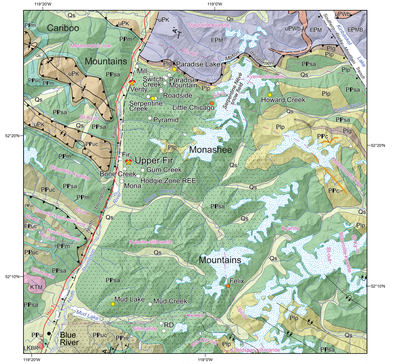


Carbonatite and related-rock complexes after Parrish & Scammell (1988), Pell (1994), Rukhlov & Bell (2010), Millonig & Groat (2013), and Rukhlov et al. (2018). Terrane geology after Nelson et al. (2013). Metamorphic and plutonic rocks after Monger & Hutchison (1971), Read et al. (1991), and Monger (2014). Morphogeological belts of the Canadian Cordillera after Gabrielse et al. (1991) and Hickin et al. (2017).

Blue River carbonatites

Carbonatites of the Blue River area underwent high-grade metamorphism and retrograde shearing during Cordilleran orogeny (Pell, 1994). They host two Ta-Nb deposits. Upper Fir is the largest and best-studied deposit, containing unusual, Ta-rich pyrochlore and ferrocolumbite, with minor fersmite and other Nb minerals, associated with locally coarse molybdenite (Rukhlov et al., 2018).

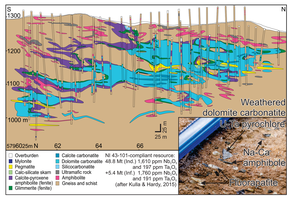
Carbonatites and related rocks of the Blue River area



Carbonatite complexes after Pell (1994), Rukhlov & Bell (2010), Millonig et al. (2012, 2013), Millonig & Groat (2013), and Rukhlov et al. (2018). Geology and metamorphic isograds after Campbell (1968), Simony et al. (1980), Raeside & Simony (1983), Pell & Simony (1987), McDonough & Murphy (1990), McDonough et al. (1991a, b, 1992), Digel et al. (1998), and Murphy (2007).

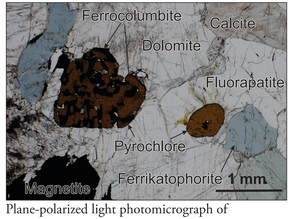
Upper Fir deposit

North-south geological cross-section through the Upper Fir carbonatite complex hosting an unusual Ta-Nb deposit, one of the largest in B.C.

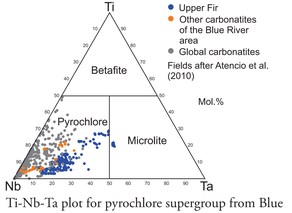


Based on 3D view of drill-hole data (Commerce Resources Corp., pers. comm.) in 150 m wide corridor. Lithological logs are shown only for a 10 m wide corridor for clarity (after Rukhlov et al., 2018).

Example of Ta-Nb mineralization in carbonatites at Upper Fir



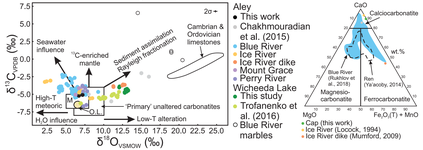
Plane-polarized light photomicrograph of granoblastic calcite-dolomite carbonatite containing Mg-Ta-rich ferrocolumbite replacing U-Ta-rich pyrochlore, Na-Ca amphibole (ferrikatophorite), fluorapatite, and Ti-bearing magnetite (after Rukhlov et al., 2018).



Ti-Nb-Ta plot for pyrochlore supergroup from Blue River carbonatites (after Rukhlov et al., 2018). Upper Fir pyrochlores are unusually Ta-rich.

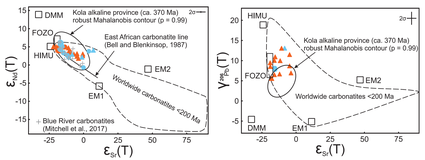
Geochemistry

$\delta^{18}O_{VSMOW}$ (‰) vs $\delta^{13}C_{VPDB}$ (‰) plot for separated calcite and dolomite from British Columbia carbonatites, related rocks, and sedimentary carbonates and CaO-MgO-[Fe_2O_3(T)+MnO] plot for the carbonatites



Arrows indicate processes causing shifts in $\delta^{18}O$ and $\delta^{13}C$ (after Demény et al., 2004 and references therein). Primary unaltered carbonatite field after Taylor et al. (1967), mantle (M) field after Deines (1989), fresh natrocarbonatite from Oldoinyo Lengai (O.L.) field after Keller & Hoefs (1995), and Cambrian and Ordovician limestone field after Veizer et al. (1999) and Chakhmouradian et al. (2015). Data from Blue River after Rukhlov et al. (2018). Most carbonatites retained their primary mantle signature despite high-grade metamorphism and deformation, which contrasts with the data from Aley, Wicheeda Lake, and Ice River Ba-REE-rich carbonatite dike.

Sr-Pb-Nd isotope correlation diagrams for the 360-330 Ma carbonatites and related rocks of British Columbia



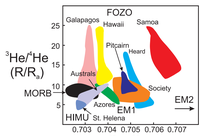
Blue River area (9 carbonatite complexes; Rukhlov et al., 2018)
Carbonatite
Alkaline silicate rock

Ice River (Locock, 1994; Mumford, 2009; Tappe & Simonetti, 2013)
Carbonatite
Alkaline silicate rock

Data from the Blue River area are for separated calcite, dolomite and apatite fractions (after Rukhlov et al., 2018). Compiled data from Ice River are for both separated mineral and whole-rock fractions (Locock, 1994; Mumford, 2009; Tappe & Simonetti, 2013). Fields for the Kola alkaline province and the <200 Ma worldwide carbonatites from Rukhlov et al. (2015). Samples having extremely radiogenic Pb isotopic compositions due to re-equilibration with coexisting U-rich minerals during metamorphism are not shown on the $\varepsilon_{Sr}(T)$ vs $\gamma_{Os}(T)$ diagram. The 360-330 Ma carbonatites and related rocks of the British Columbia alkaline province follow mixing trends involving FOZO (FOcus ZOne), HIMU (high-$^{238}U/^{204}Pb$ or μ), and EM1 (enriched mantle 1) mantle end-members found in oceanic basalts, young (<200 Ma) carbonatites worldwide, and the plume-related Kola alkaline province (ca. 370 Ma). FOZO, HIMU, EM1, and EM2 are considered to be the lower mantle components sampled by mantle plumes (e.g., Hart et al., 1992; Hofmann, 2014). Note that the depleted, mid-ocean ridge mantle (DMM) end-member, representing the upper mantle, is excluded from the mixing trends defined by the carbonatite data.

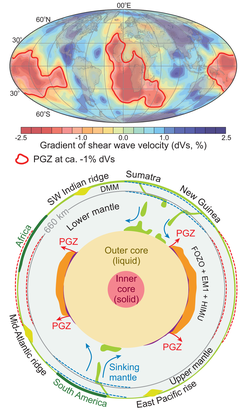
Mantle source

Sr vs He isotope correlation plot for oceanic basalts (from Keken et al., 2002)



The data from OIBs define elongated mixing arrays converging on the FOZO component at high $^3He/^4He$ ratios, indicating relatively undegassed, i.e. deep-seated source of the hotspot lavas.

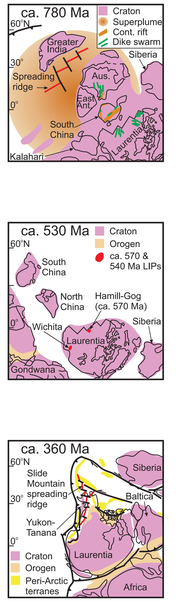
Shear wave velocity at the core-mantle boundary and equatorial cross-section (modified from Burke, 2011)



Plume generation zones (PGZ) outline two large low shear wave velocity provinces (LLSVP) at the core-mantle boundary (D"), which were stable in the present position for at least hundreds of m.y. (Burke, 2011).

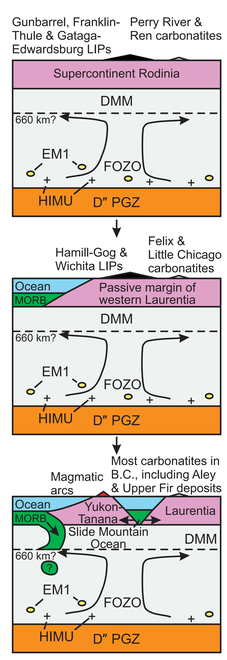
Paleogeography

Reconstructions of Neoproterozoic to Late Paleozoic continents (modified from Li et al., 2008; Ernst & Bleeker, 2010; Nelson et al., 2013)



Paleogeographic reconstructions place the western margin of Laurentia, and hence the 810-330 Ma British Columbia alkaline province, above the equatorial LLSVP marked by a long-lived plume-generation zone (PGZ) at the core-mantle boundary. Therefore, we suggest that parental magmas of carbonatites and related rocks of the British Columbia alkaline province episodically tapped a long-lived, deep-mantle plume from the PGZ during a period of ca. 480 Ma since the Neoproterozoic.

Summary model



ca. 810-700 Ma

Protracted breakup of the supercontinent Rodinia due to arrival of a deep-mantle plume (e.g., Li et al., 2008).

ca. 500 Ma

Passive margin development on the western flank of Laurentia (Bond & Kominz, 1984; Ross, 1991; Colpron et al., 2002).

ca. 360-330 Ma

Rifting the continental margin and initiating the Slide Mountain ocean as a back-arc basin (e.g., Nelson et al., 2013).

Conclusions

British Columbia carbonatites are indistinguishable from worldwide carbonatites generated by deep-mantle plumes. We suggest that the Cordilleran examples were derived from a long-lived, deep-level mantle plume that was tapped episodically since the Neoproterozoic.

Acknowledgments

This study was supported by British Columbia Geological Survey. Isotopic analyses were carried out at the Pacific Centre for Isotopic and Geochemical Research. We thank G. Fortin (British Columbia Geological Survey) for assisting during the 2018 fieldwork, and T.D. Peterson (Geological Survey of Canada) and J.A. Brown (TerraLogic Exploration Services) for donating samples of the Ice River carbonatites.

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