# Composition of Logtung Beryl (aquamarine) by ICPES/MS: A Comparison with Beryl Worldwide 

by Mitchell G. Mihalynuk and Ray Lett

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## INTRODUCTION

Beryl is a beryllium-aluminum-silicate mineral. When it grows unimpeded, it forms hexagonal prisms. In its simplest form, its chemical formula is $\mathrm{Be}_{3} \mathrm{Al}_{2}\left[\mathrm{Si}_{6} \mathrm{O}_{18}\right]$, and it is clear or white in colour. Addition of $\mathrm{Cr}, \mathrm{V}, \mathrm{Mn}$, or Fe into the formula can result in some of the world's most sought after gemstones: green emerald $(\mathrm{Cr}, \pm \mathrm{V})$, pink morganite $(\mathrm{Mn})$, yellow heliodor ( Fe ), and blue aquamarine ( Fe ). Yet, no simple addition of these key elements will do; the amount and valence of the element and site of its occupancy within the crystal lattice are key to colour. Beryl's molecular structure is dominated by hexagonal rings comprised of six Si-O tetrahedra (Figure 1). Linking the rings are tetrahedrally coordinated $\mathrm{Be}( \pm \mathrm{Fe}, \mathrm{Mn}, \mathrm{Mg})$ and octahedrally coordinated $\mathrm{Al}( \pm \mathrm{Fe}, \mathrm{Cr}, \mathrm{V})$. The rings are stacked parallel to the crystal prism (z axis) producing hollow channels. This "zeolite" structure, named after a family of minerals known as molecular sieves, permits occupancy of water and cations (particularly alkalis, $\mathrm{Na}, \mathrm{Li}, \mathrm{K}, \mathrm{Cs}, \mathrm{Rb}$ ) within the crystal lattice, but outside of the crystal formula unit. Thus, up to $8 \%$ alkalis (Deer et al., 1966) in igneous beryl, and an amazing $14 \%$ in hydrothermal beryl (Vianna et al., 2002b) can be accommodated along with a charge balance not accounted for in the formula unit. Schaller et al. (1962) described beryl chemistry as an isomorphic series between the end members $\mathrm{Be}_{3} \mathrm{Al}_{2} \mathrm{Si}_{6} \mathrm{O}_{18}$, ( $\mathrm{Na}, \mathrm{Cs}$ ) $\mathrm{Be}_{2} \mathrm{Al}(\mathrm{Al}$, $\mathrm{Li}) \mathrm{Si}_{6} \mathrm{O}_{18}$, and $(\mathrm{Na}, \mathrm{K}, \mathrm{Cs}) \mathrm{Be}_{3} \mathrm{R}^{3+} \mathrm{R}^{2+} \mathrm{Si}_{6} \mathrm{O}_{18}$ where $\mathrm{R}^{3+}=\mathrm{Al}$, $\mathrm{Fe}, \mathrm{Cr}$ and Sc , and $\mathrm{R}^{2+}=\mathrm{Be}, \mathrm{Fe}, \mathrm{Mn}$ and Mg . In consideration of the channel components, the formula may be generally presented as: $\mathrm{R}^{2+} \mathrm{Be}_{2} \mathrm{R}^{3+} \mathrm{Al}\left[\mathrm{Si}_{6} \mathrm{O}_{18}\right] .\left(\mathrm{H}_{2} \mathrm{O}\right)(\mathrm{Na}, \square)$.

Refractive index of beryl increases with increasing content of alkalis in the channels as the channels would otherwise be empty space within the crystal structure. When iron occupies the channels it appears to be a quintessential agent in the colouration (i.e. the chromophore) of aquamarine gemstones (Price et al., 1976; Vianna et al., 2002a). A correlation between Fe content and colour of aquamarine has long been known, although some authors have attributed colouration to other agents such as Cs (Sosedko, 1957). Mössbauer spectroscopy in particular has helped to isolate the chromophore's position within the crystal lattice. Combining this technique with other spectroscopic methods it can be reasonably inferred that ferric iron with octahedral coordination produces a yellow colour, while
ferrous iron in the octahedral site has no effect on colour, and ferrous iron within channels produces a deep blue colour (Vianna et al., 2002b). These same authors showed that in one small population of samples deep blue aquamarines have little $\mathrm{Fe}^{3+}$ and greener samples contain more $\mathrm{Fe}^{3+}$ or less $\mathrm{Fe}^{2+}$ in channels. The common process of heat-treating aquamarine to produce darker blue colour may work through charge transfer of $\mathrm{Fe}^{3+}$ to $\mathrm{Fe}^{2+}$, thereby eliminating the greenish cast (Blak et al., 1982).

In this paper we report on the chemical composition of blue beryl that is abundant at the Logtung tung-sten-molybdenum deposit of northwestern BC (Figure 2). Mineralization is associated with a porphyritic quartz monzonite intrusion dated at $\sim 58 \mathrm{Ma}$ ( $\mathrm{U}-\mathrm{Pb}$ zircon; Mihalynuk and Heaman, 2002) although previous isotopic age determinations suggest an age of around $109 \pm 2 \mathrm{Ma}$ (K-Ar muscovite; Hunt and Roddick, 1987) to $118 \pm 2 \mathrm{Ma}$ ( $\mathrm{Rb}-\mathrm{Sr}$; Stewart and Evensen, 1983).Observations reported on here are limited to a few person-days devoted mainly by Ministry of Energy and Mines personnel to mapping the area around the Logtung deposit as part of the National Mapping Program (Ancient Pacific Margin NATMAP, 1999-2000; Mihalynuk et al., 2000). Geological work un-


Figure 1. Atomic structure of Beryl, modified after (Deer et al., 1966). Atom loci are projected onto basal 0001 plane and the lower Si-O ring is shown by dashed bonds. Grey trapezoid outlines the unit cell. Ring channels may be occupied by water alkalis and ferrous iron.


Figure 2. Location of the Logtung porphyry, adapted from (Mihalynuk and Heaman, 2002).
der the NATMAP program was not specifically aimed at beryl, however, blue beryl was reported previously at the Logtung deposit (MINFILE number 1040 016) and we found to it be common, with crystals up to 1.5 cm in diameter and 3 cm long. Unfortunately, all large specimens observed were cloudy or heavily included. Clear crystals tended to be less than 3 mm in diameter.

If large, gem quality stones are to be found at Logtung, a more methodical and intense exploration program than our ancillary approach will be required. To help determine if such exploration is worthwhile, we analyzed a sample of Logtung beryl to compare its chemical affinity with known gem occurrences from around the globe. The chemistry and structure of beryl from some of these occurrences is very well characterized, providing a good benchmark for comparison.

## Methods

About 2 grams of beryl were extracted from a single fist-sized sample, crushed with an agate mortar and pestle to a particle size of less than $\sim 1 \mathrm{~mm}$ and then carefully sorted. Clear, inclusion-free grains were separated and powdered in the mortar and pestle, resulting in a final sample yield of 0.8 g . The powder was submitted to ACME Analytical Laboratories, Vancouver, for analysis by lithium metaborate $\left(\mathrm{LiBO}_{2}\right)$ fusion followed by Inductively Coupled Plasma Emission and Mass Spectrometry (ICPES/MS).

A sample aliquot of 0.2 g was fused at $900^{\circ} \mathrm{C}$ with $\mathrm{LiBO}_{2}$ which enables complete dissolution of the residue in $\mathrm{HNO}_{3}$ for analysis by ICPES/MS. This technique provides an analytical sensitivity with a small sample size that is comparable to wet chemical or microprobe analyses of beryl published previously (Table 1). Benefits of ICPES/MS include ease of sample preparation, and the large number of major and trace elements can be determined simultaneously. Drawbacks include the inability to analyze $\mathrm{Li}_{2} \mathrm{O}$ content because of the addition $\mathrm{LiBO}_{2}$ and possible loss of Li during fusion, inability to distinguish Fe valance state, and lack of volatiles determination (e.g. $\mathrm{H}_{2} \mathrm{O}$ ). As discussed above, $\mathrm{H}_{2} \mathrm{O}+$ and $\mathrm{H}_{2} \mathrm{O}$ - are principally restricted to crystal channels and do not enter into the crystal structure. $\mathrm{Fe}^{3+} / \mathrm{Fe}^{2+}$ can be estimated from cationic site occupancy balance. However, lack of $\mathrm{Li}_{2} \mathrm{O}$ is a serious drawback because the beryl crystal structure can incorporate a large proportion of Li. In the case of the Logtung beryl, Li may be an important constituent based upon the low oxide totals and the results of cationic ratio calculations (Tables 1 and 2).

## Results

ICPES/MS analytical results are reported at the bottom of Table 1. ICPES/MS detection limits reported are comparable or better than those reported for the results by other methods listed in Table 1. Figure 3a shows that the most abundant oxides of Si and Al are low in the Logtung beryl
TABLE 1．COMPOSITION OF BERYL FROM OCCURRENCES WORLDWIDE．





|  | 0.01 | 0.04 | 0.14 |
| ---: | ---: | ---: | ---: |
| $\mathbf{0 . 1}$ | $\mathbf{0 . 0 1 4}$ | $\mathbf{0 . 2 0 9}$ |  |
| 04 | 0.006 | 0.001 |  |


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| U゙ | べ | $\stackrel{\square}{0}$ | $00000000000$ |








| Sample No． | Ref． | Method＊＊ |
| :---: | :---: | :--- |
| 1．Ingersoll | 1 |  |
| 2．Aquamarine | 1 |  |
| 3．Biotite schist | 1 |  |
| Be 609 | 2 |  |
| Be 522 | 2 |  |
| Be551 | 2 |  |
| DHZ4 | 3 |  |
| 1 | 4 | microprobe\＃ |
| 2 | 4 | microprobe\＃ |
| 3 | 4 | microprobe\＃ |
| 4 | 4 | microprobe\＃ |
| DHZ1 | $5^{*}$ |  |
| DHZ2 | $6^{*}$ |  |
| DHZ3 | 7 |  |
| Rbba1 | 8 | XRF／EDS／ICP |
| Rbba2 | 8 | XRF／EDS／ICP |
| Rbip1 | 8 | XRF／EDS／ICP |
| Rbip2 | 8 | XRF／EDS／ICP |
| Rblf1 | 8 | XRF／EDS／ICP |
| Rblf2 | 8 | XRF／EDS／ICP |
| Rbsc1 | 8 | XRF／EDS／ICP |
| Rbsc2 | 9 | XRF／EDS／ICP |
| MMI03－Be | 9 | ICPES／MS |
| ICPES／MS detection limit |  |  | ＊$=$ in Deer，et al．（1966） References（Ref．）： $1=$ Jacob，J．，1938； $2=$ Radcliffe and Campbell，1966； $3=$ Schaller，W．T．，et al．，

Seelye，F．T．，1945； $6=$ Quesnel，P．，1937； $7=$ Sosedko，T．A．，1957； $8=$ Schaller，W．T．et al．， 1962 $* *=$ blank if not reported，incompletely reported，or if a wide variety of wet chemical and spectrographic methods employed on an element by element basis．
Mean $\mathrm{SiO}_{2}$ and $\mathrm{Al}_{2} \mathrm{O}_{3}$ excluding Logtung analysis are： 65.12 and $17.7 \mathrm{Wt} \%$ respectively．
\＃－Li $\mathrm{Li}_{2} \mathrm{O}, \mathrm{BeO}$ and $\mathrm{H}_{2} \mathrm{O}$ by standard wet chemical techniques；identical values for some oxides of samples 3 and 4 indicate not analyzed but averaged 1 and 4 ．

| EII0 | Stor | 269 乙 | ts801 |
| :---: | :---: | :---: | :---: |
| L90＇0 | LE6 ${ }^{\text {I }}$ | 626 て | 601 |
| 98て＇0 | $618{ }^{\text {I }}$ | £90＇$\varepsilon$ | tLI＇II |
| 8Lで0 | 2E8＇I | ILで\＆ | ¢8E＇II |
| £010 | 8t6 ${ }^{\text {I }}$ | $8 \pm 6$ て | 200＇II |
| $960{ }^{\circ}$ | t96．${ }^{\text {I }}$ | ちて6て | £6601 |
| zzoº | 2z6 ${ }^{\text {I }}$ | LE0＇$\varepsilon$ | 98601 |
| †200 | $696{ }^{\text {I }}$ | 200＇$\varepsilon$ | II |
| It900 | Z66 I | 6＋8＇て | 62S＇II |
| $16 \mathrm{Z}^{\circ}$ | $90^{\circ}$ | \＆ 28.2 | tLI＇II |
| LLO＇0 | $880{ }^{\circ}$ | 958＇z | szo＇il |
| LZOO | 800 Z | £00＇$\varepsilon$ | 880 ${ }^{\circ} \mathrm{I}$ |
| szo 0 | $666{ }^{\text {I }}$ | z 20 ¢ | 970 －It |
| £0\％ | $696{ }^{\text {I }}$ | $850 \cdot \varepsilon$ | Lto ${ }^{\text {a }}$ |
| $810^{\circ}$ | ¢¢0＇z | $186{ }^{\text {® }}$ | t¢0 ${ }^{\text {II }}$ |
| ¢Es ${ }^{\circ}$ | 2t6 I | 620 ¢ | LtS＇II |
| $690{ }^{\circ}$ | て9t＇z | 69 でて | E08．01 |
| $6+00$ | \＆とを゙て | ¢z9｀て | 86601 |
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| Lt9 0 | ¢S0＇て | 985＇て | $8 \mathrm{Sc}^{\prime} \mathrm{I}$ |
| 8てz゙0 | $986{ }^{\text {I }}$ | $668{ }^{\circ} \mathrm{Z}$ | عıl＇ı |
| してで0 | tSO＇z | L18＇て | $8 \mathrm{IJ}^{\circ}$ |
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Colour
clear
blue
green
green
yellow
aqua
bluish
med．blue
light blue
light green
med．green
pale green
clear
pink
green－blue
green－blue
light blue
light blue
dark blue
dark blue
light blue
light blue
 MMI03－Be Logtung，BC
 $\mathrm{Fe} 2 \mathrm{O} 3+\mathrm{V} 2 \mathrm{O} 3$ ．＂ $\mathrm{R}+$＂is estimated total cations in channels calculated as： $\mathrm{Na}+\mathrm{K}+$ excess FeO


Legend
$\times$ colourless
yellow
pink
dark green
$\Delta$ light green
- greenish blue
- light blue
(1) (Logtung)
- medium blue
- dark blue


Figure 3. A. variance of $\mathrm{SiO}_{2}$ and $\mathrm{Al}_{2} \mathrm{O}_{3}$ content of beryl about a visually estimated mode from analyses in Table 1 of $66 \mathrm{wt} . \% \mathrm{SiO}_{2}$ and 18 wt. $\% \mathrm{Al}_{2} \mathrm{O}_{3}$. Logtung analysis is highlighted by a shaded circle.
when compared to the visually estimated mean of $65 \mathrm{wt} . \%$ $\mathrm{SiO}_{2}$ and $18 \mathrm{wt} . \% \mathrm{Al}_{2} \mathrm{O}_{3}$ from the data set, possibly contributing to the low total and the 0.3 atomic formula unit deficit in the tetrahedral site (Th, Table 2).

Iron on the other hand, compares favourably with Fe contents in blue (and green) gem beryl worldwide (Figure 3b, c). Only beryl which lack blue or green colouration have lower Fe contents as listed in Table 1 and shown on Figures 3b and 3c.

## Conclusions

Inductively Coupled Plasma Emission and Mass Spectrometry (ICPES/MS) permits analysis of small samples $(0.2 \mathrm{~g})$ with low detection limits. As a result, the method is ideally suited to elemental determination of coarse, single crystals. Drawbacks of the method are acute for minerals that contain significant $\mathrm{Li}_{2} \mathrm{O}$ because fusion of the sample for ICPES/MS analysis requires addition of $\mathrm{LiBO}_{2}$. Where OH - is a key components of the crystal chemistry, it must be also be determined by a separate method. Nevertheless,

ICPES/MS provides a quick and cost effective method for determining the elemental abundance of key chromophores. In the case of the Logtung beryl, Fe content compares with gem aquamarine (and emerald) elsewhere in the world. On the basis of Fe content, the Logtung beryl can be classified as aquamarine, although the presence of gem-quality aquamarine at this locality has yet to be established.

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